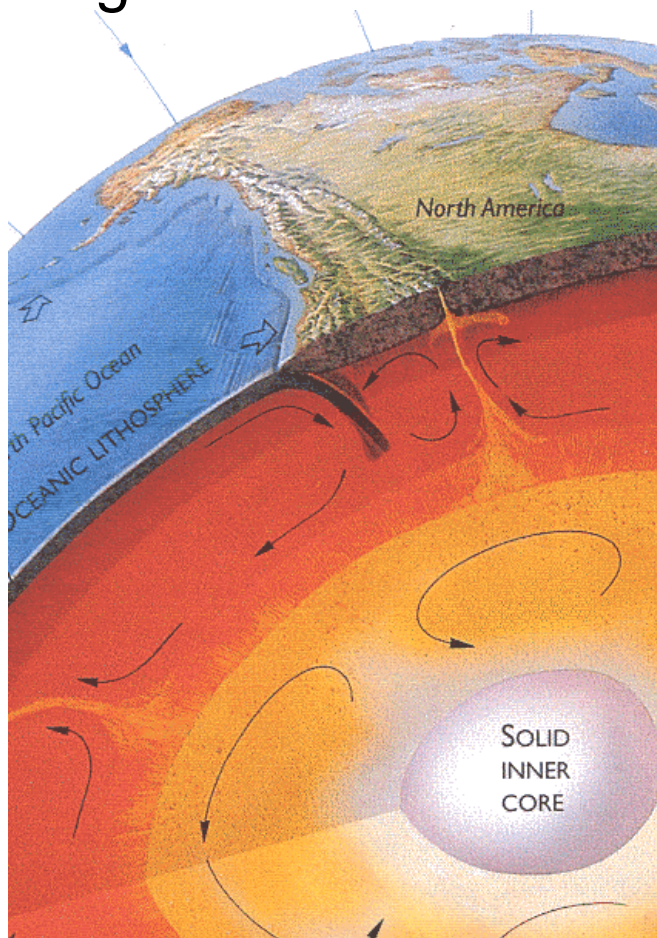


# Significance of seismic tomography within the wider geoscience community

- Computational seismology works on improving imaging methods. Other geoscientists care mainly about the images themselves.
- Hopefully, tomography results make sense in light of surface studies (geology, tectonics) and mantle convection simulations (geodynamics).
- Solutions are non-unique -- but how bad is it? And how do I **convey** that to a non-specialist?

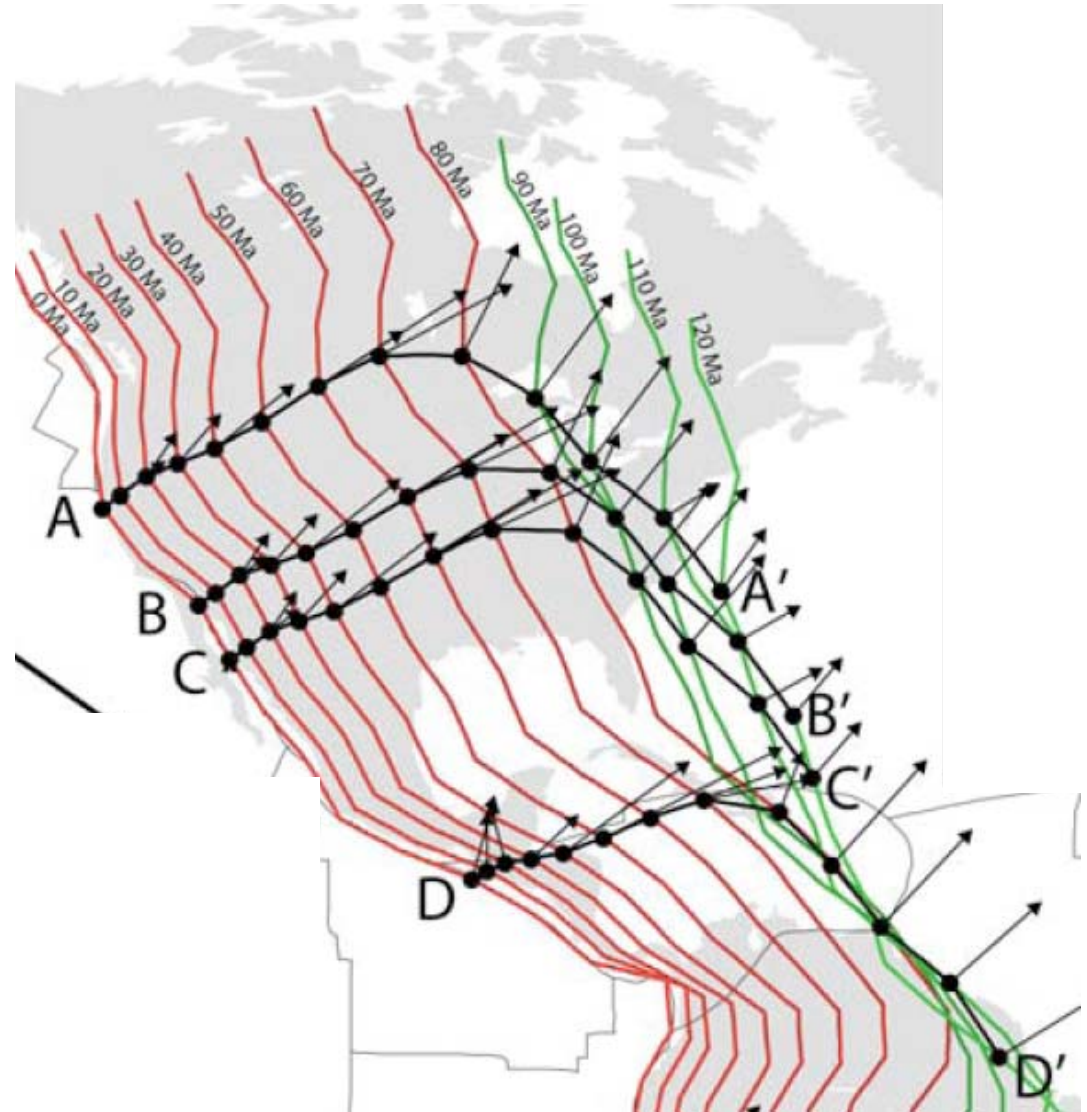
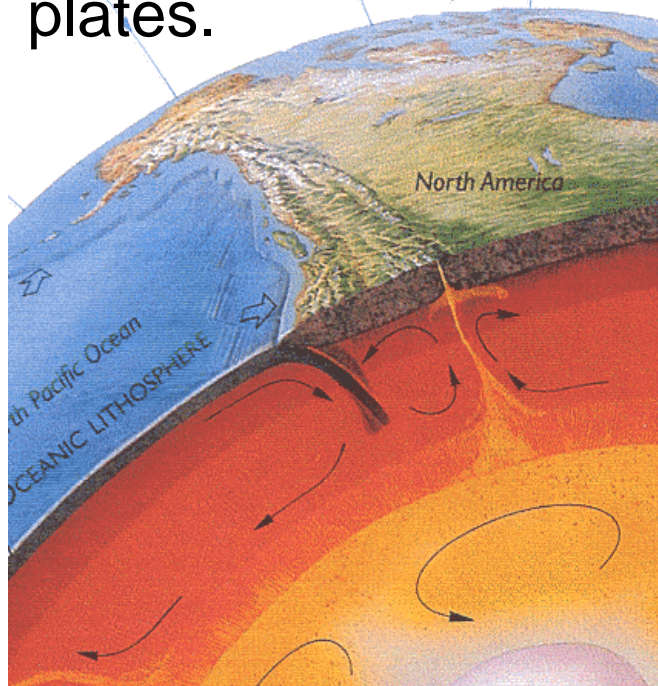
# Example study: the mantle under North America -- window into a distant past

Subduction on the west coast: an oceanic plate gets submerged beneath the continent



# A simple tectonic history(?)

- A single large plate has been subducting beneath the west coast for 180 million years. No significant interference from other plates.

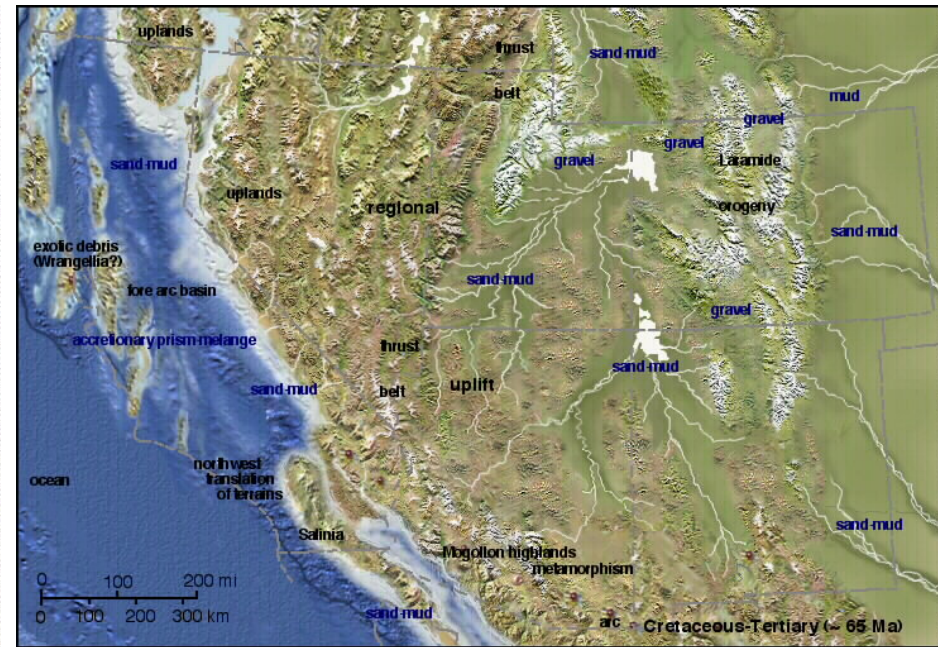




# Geologists puzzle about episodes of rapid change in the past

Western U.S., 75 Myr ago

65 Myr ago



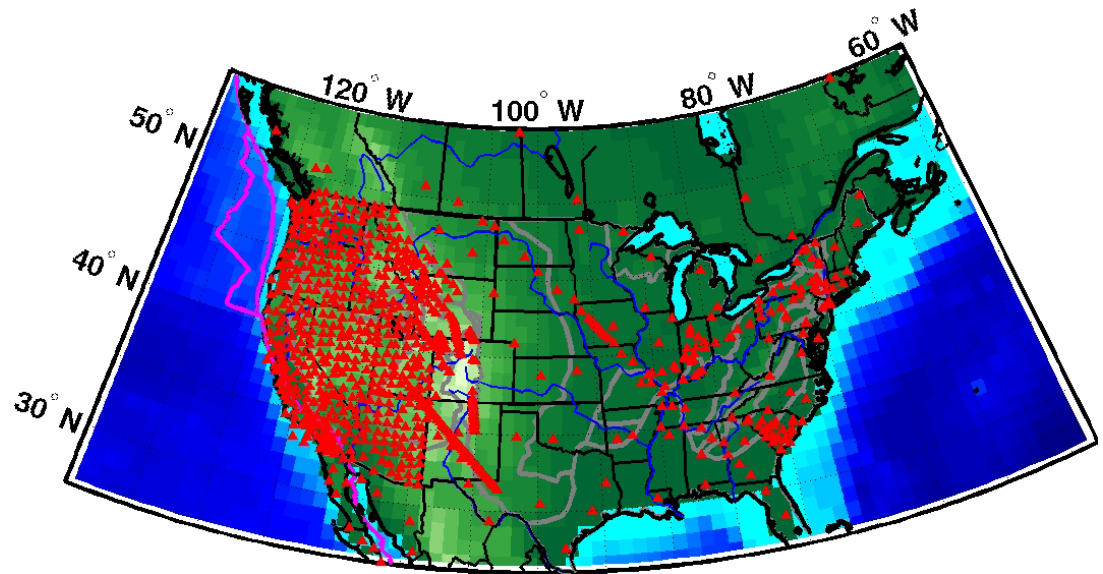
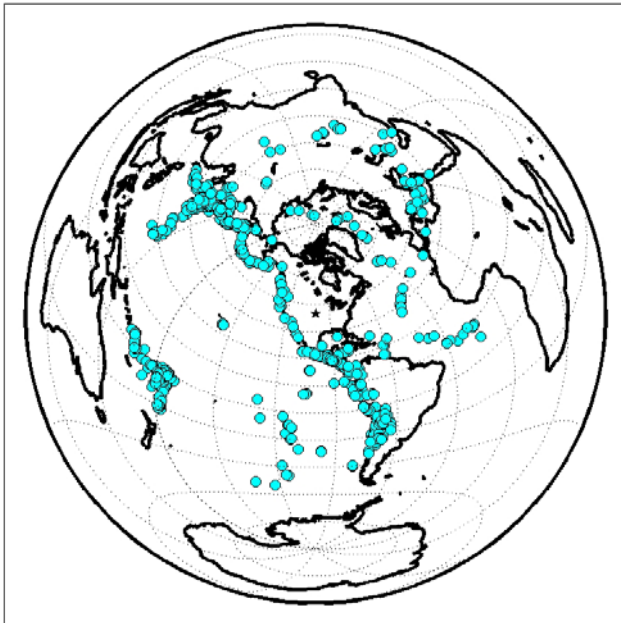
75 million years ago: A shallow inland sea covers the Rocky Mountain area

5-10 million years later: the sea is gone; 4-km-high mountains have risen.

# Geometry of the body-wave tomography experiment

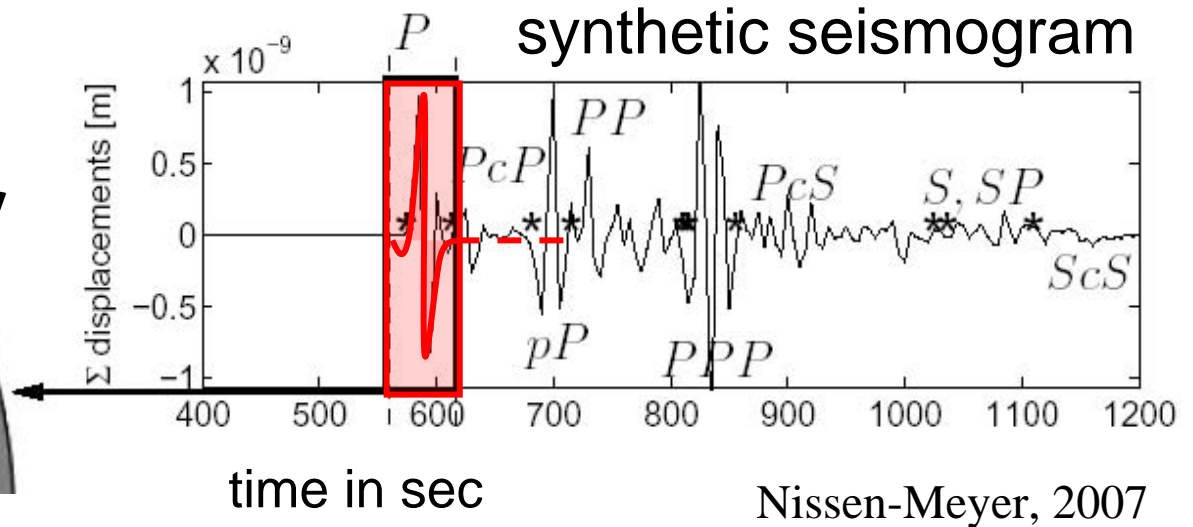
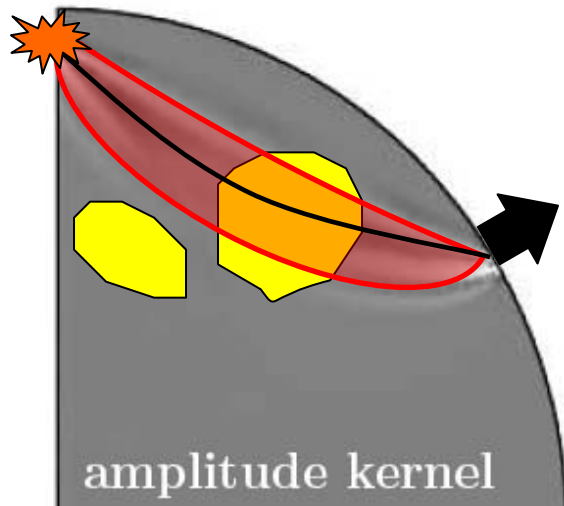
637 earthquake sources

1125 broadband receivers (seismometers)

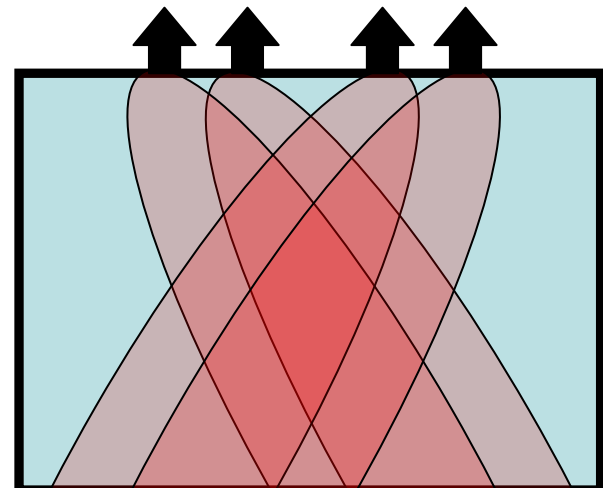


- Signal sources are P-waves generated by large but distant earthquakes

# How body waves sample the earth

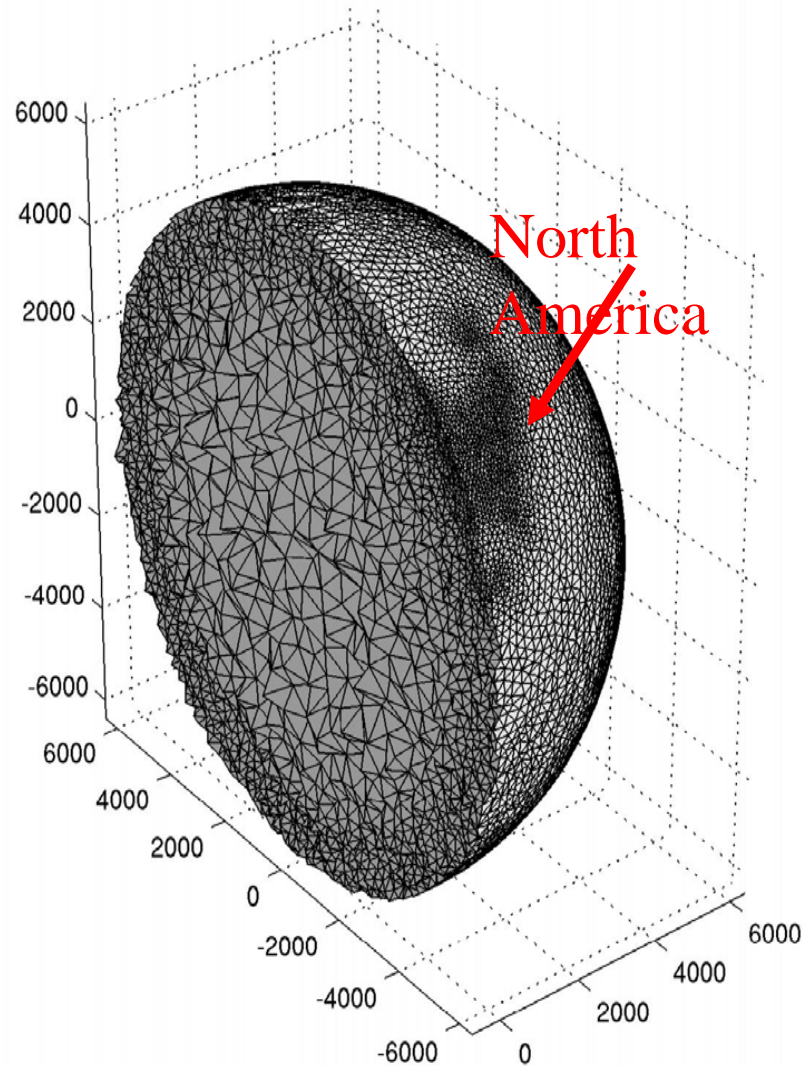


Sensitivity kernels (shaded red) map the areas sampled by the body waves used.





# To solve the inverse problem, expand the sensitivity kernels on a global grid



# Inverse problem (linear)

$$\begin{array}{c} \text{dT} \\ \text{reg} \end{array} \begin{array}{c} \text{=} \\ \text{=} \end{array} \begin{array}{c} \mathbf{K}_T \\ \text{reg} \end{array} * \text{dln}(v_p)$$

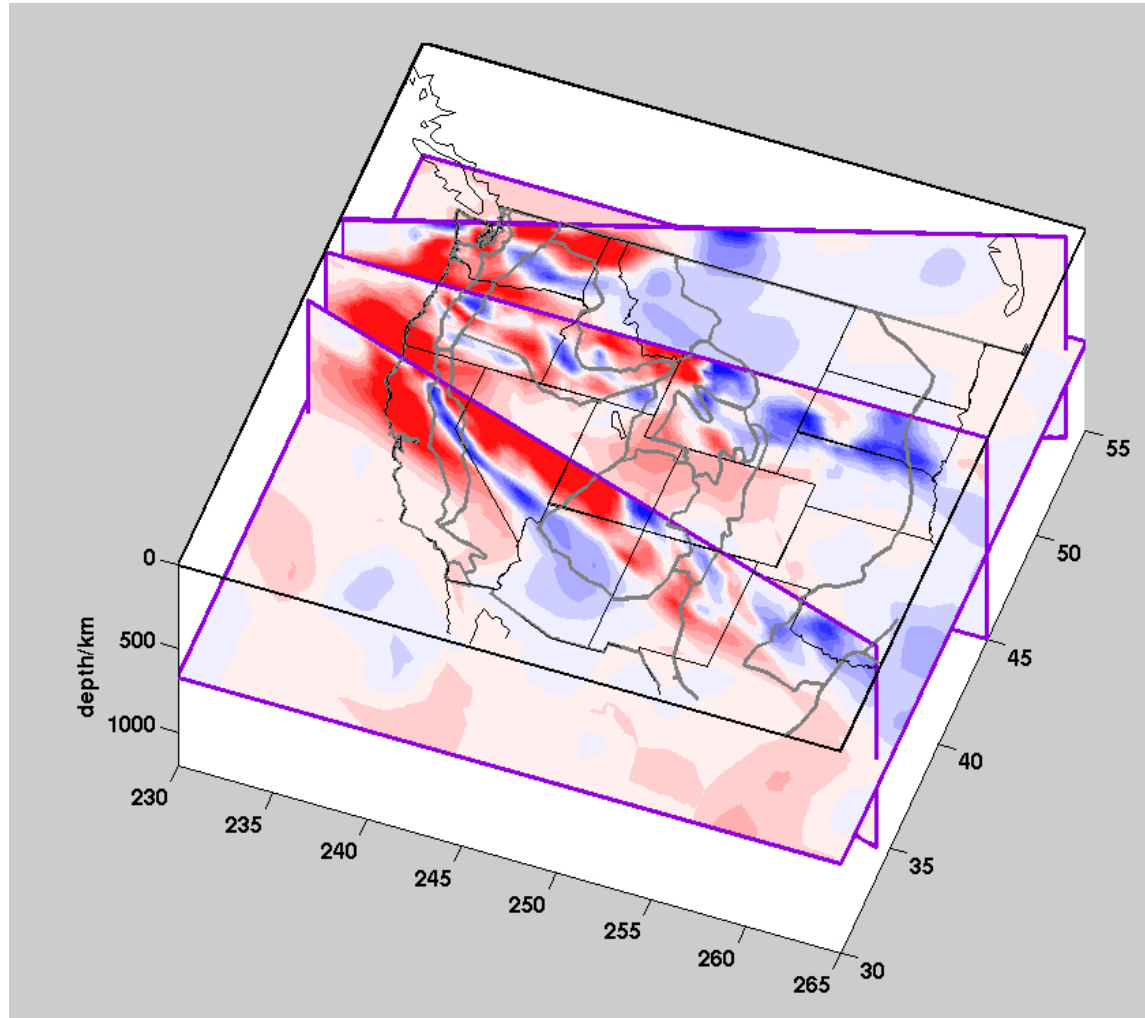
data

physical model  
(rows = kernels)

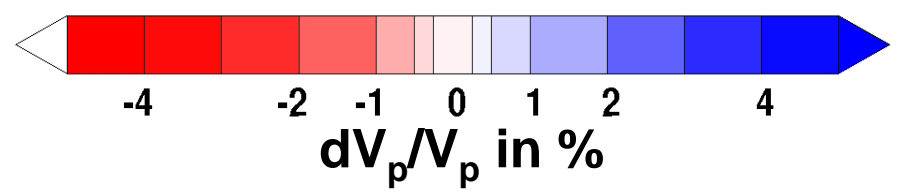
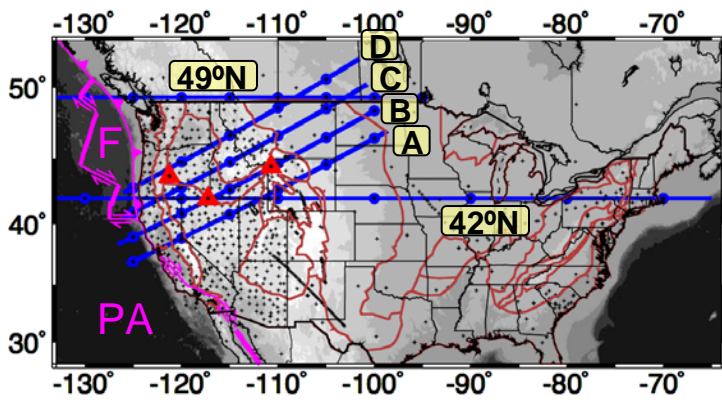
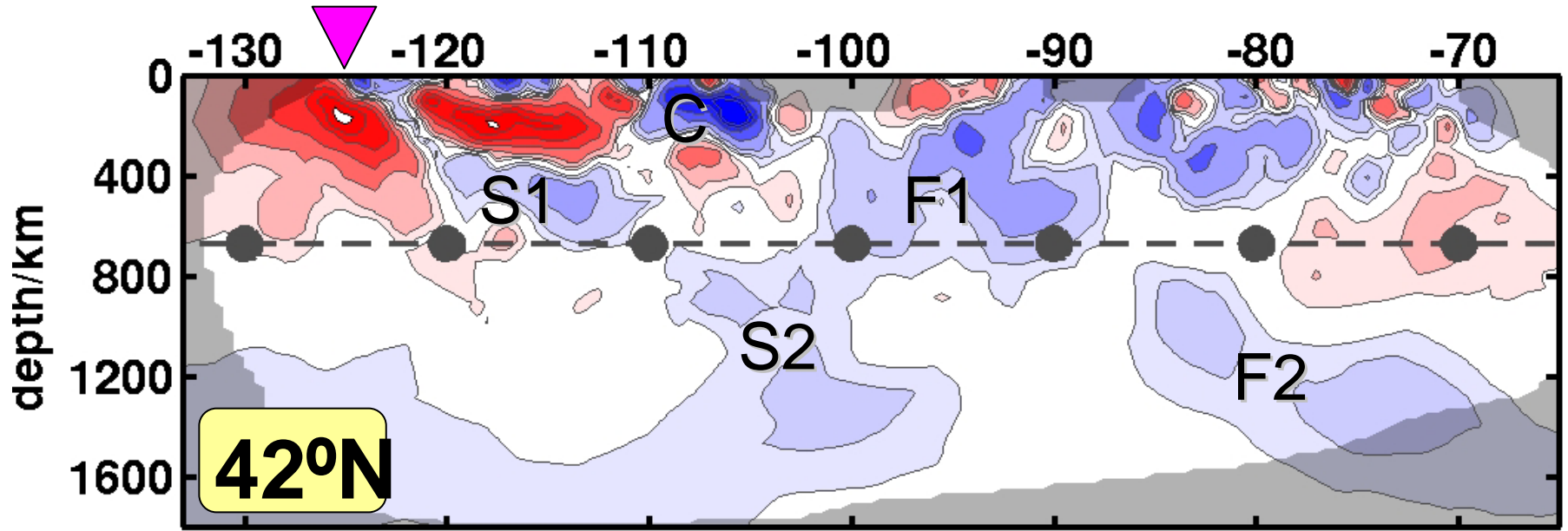
3-D earth structure  
(unknowns)



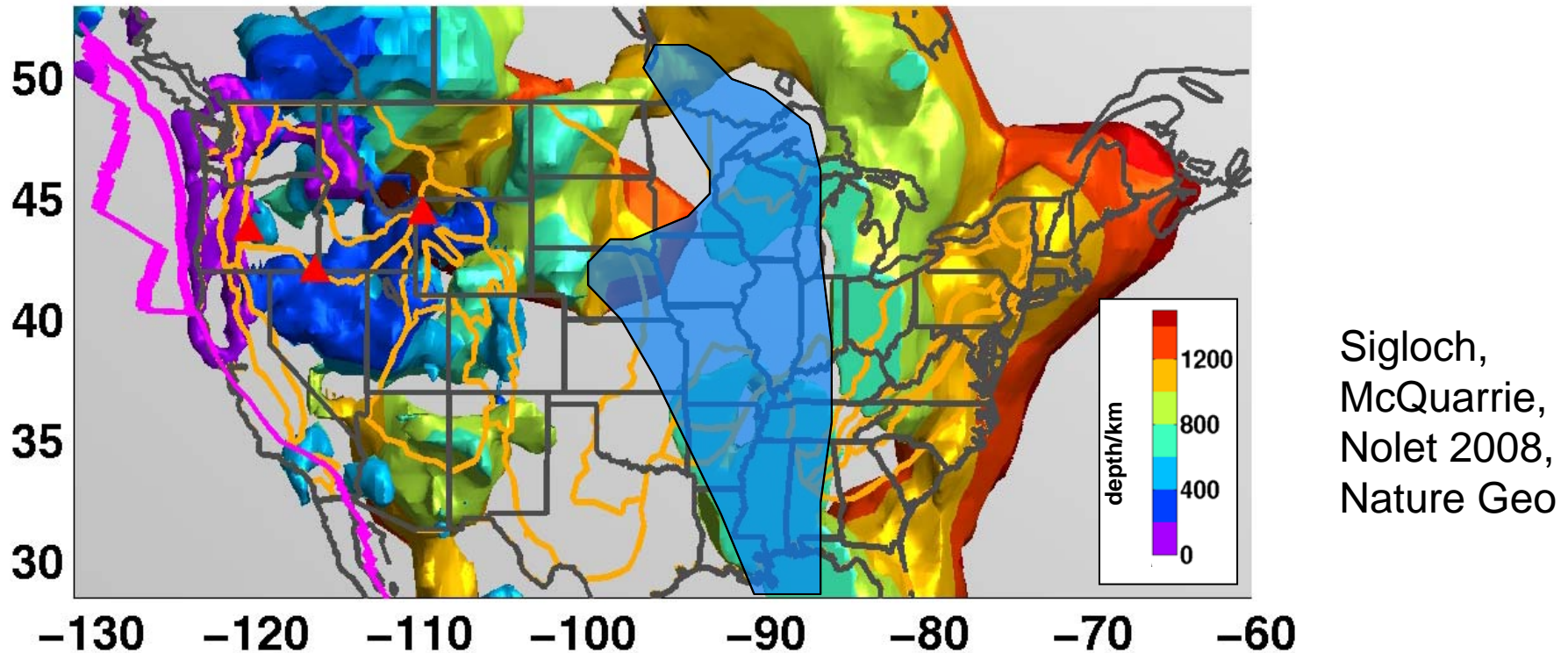
# Result: a 3-D model of seismic P-wave velocities under North America



The subducted slab (blue=fast=cold) in the mantle down to 1800 km depth



# Image of the subducted Farallon slab in the mantle

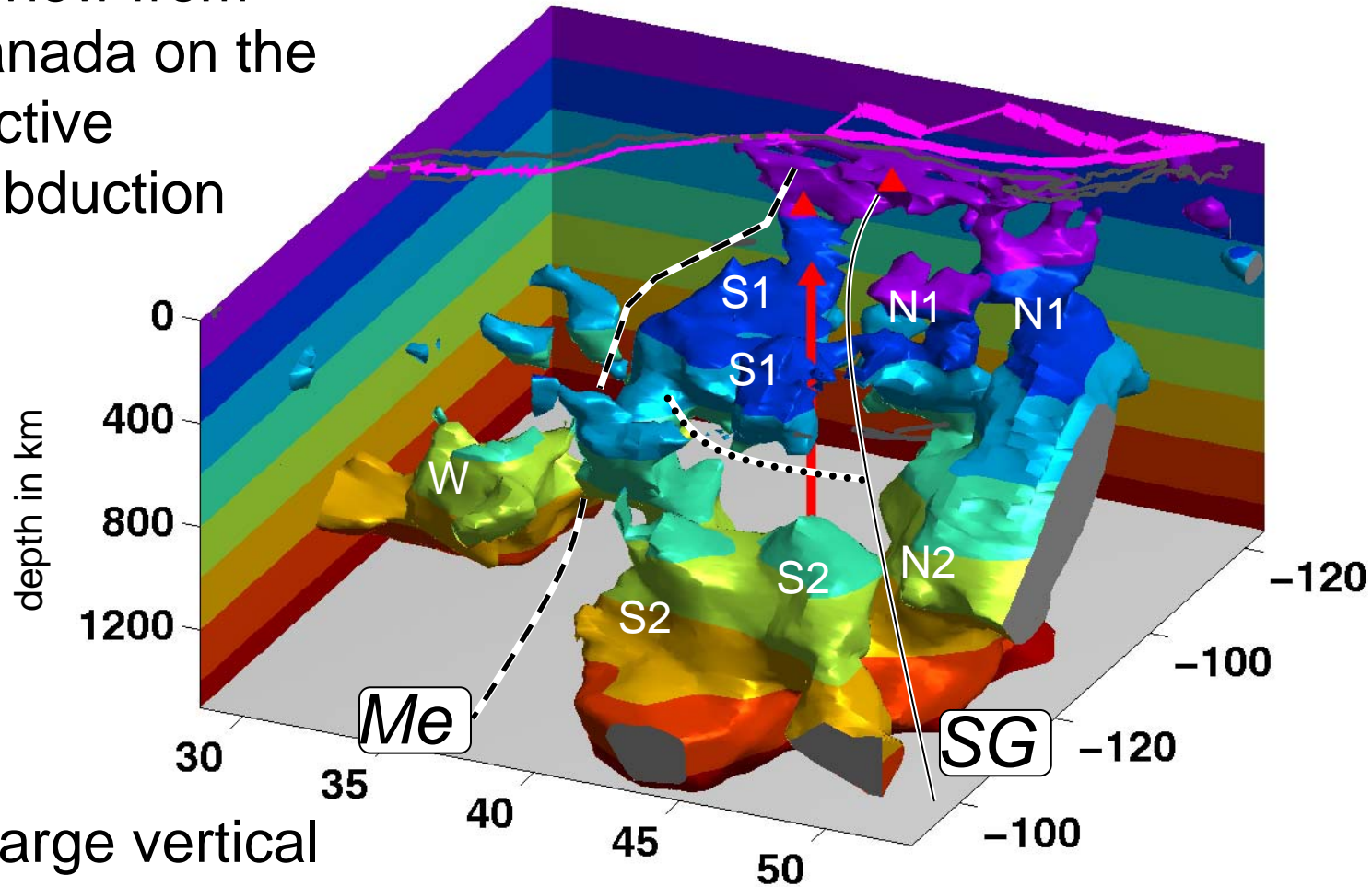


- Seismically fast material is contoured (fast means cold).
- Color signifies depth. We can confidently image ~1500 km deep.
- Crust and lithosphere not rendered.



# The “current” subduction system

Bird’s-eye view from eastern Canada on the currently active western subduction system.

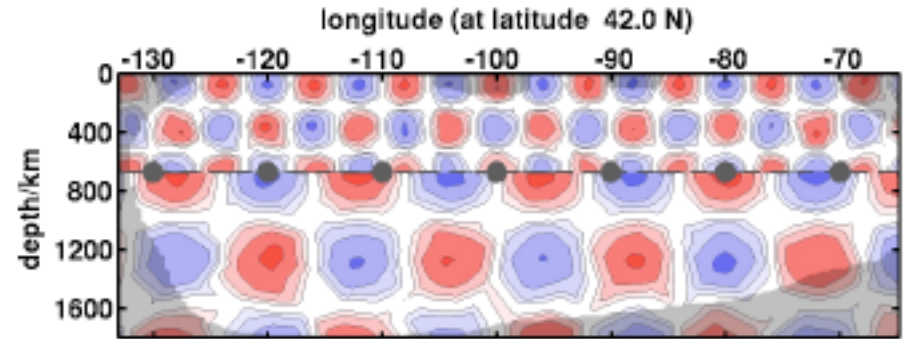


Notice the large vertical offsets of some adjacent fragments.

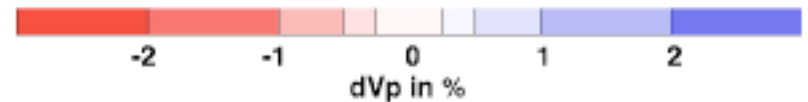
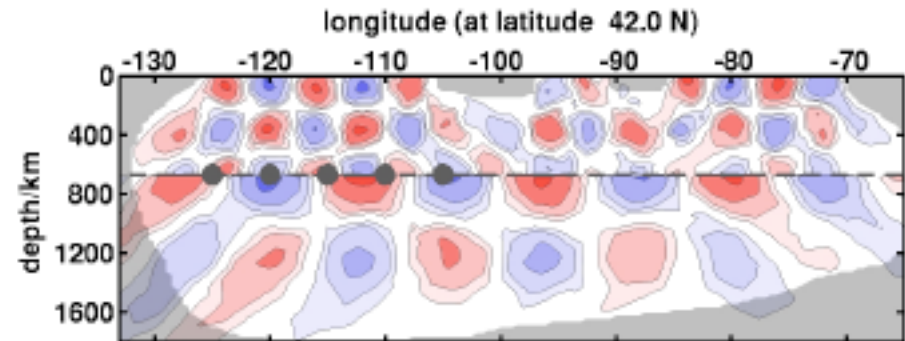
# Resolution tests

- Assume some hypothetical earth structure
- Compute synthetic data from it
- Invert the data. Is the input structure recovered?

INPUT



RECOVERED



# Questions

- New model is surprising but plausible (because it explains more geological observations than earlier models).
- How different can other plausible models be that also fit the data? Possible to generate such models a priori?
- Alternatively, at what level of confidence can we say that certain **interesting** features are real? (Example: tears in the plate, which are geodynamically important.)



# Inverse problem

