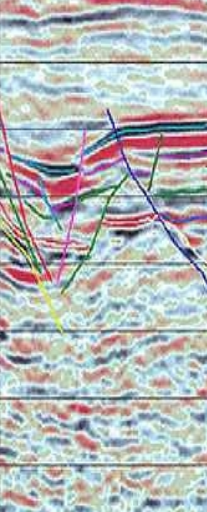



Seismic Tomography: Example of a geophysical inverse problem



Dr. Karin Sigloch (karin.sigloch@lmu.de)
Theresienstr. 41, Zi. 445

1. What is an inverse problem?
 2. A real-world example: tomography of the Earth's mantle under North America
 3. How does it work?
- 



Q: What is an „inverse problem“?



A: An indirect measurement.

We want to measure some important „EARTH_PROPERTY“ (e.g., seismic velocity $v(x)$), and have no tools to do it. Instead we know how to measure some other property called „DATA“ (e.g., traveltimes dT) And we know some phys./math. relationship „MAPPING_FCT“, so that:

$$\text{DATA} = \text{MAPPING_FCT}(\text{EARTH_PROPERTY})$$



If we are able to find an „inverse function“ MAPPING_FCT^{-1} so that

$$\text{EARTH_PROPERTY} = \text{MAPPING_FCT}^{-1}(\text{DATA}),$$

then the problem is solved.

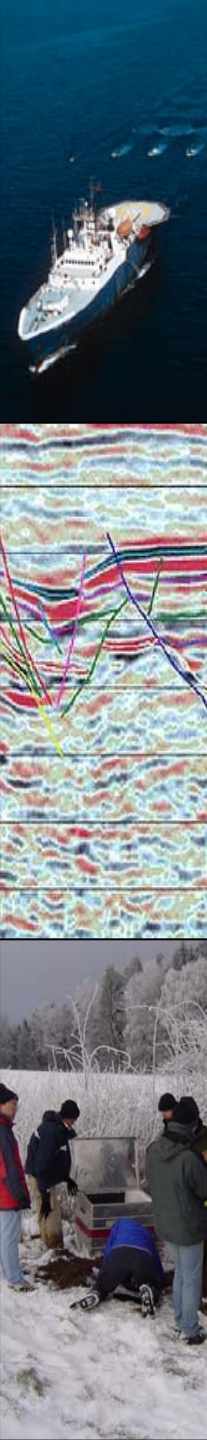
Inverse problems are common

Seismology

EARTH_PROPERTY: as a function of space (x,y,z) ,
e.g., P -velocity or intrinsic attenuation, or rock
composition

DATA: Seismograms (and data derived from them,
like traveltimes, amplitudes...) at discrete points
at the surface

MAPPING_FCT: wave equation (or some
approximation to it, like rays from Snell's law)



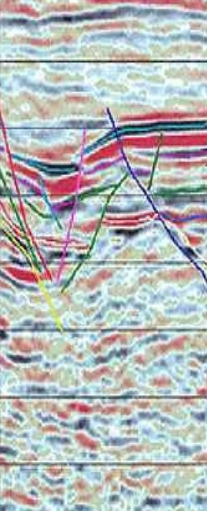
Inverse problems are common

Medical Imaging: Computed Tomography

EARTH_PROPERTY: structure of tissue in the human body

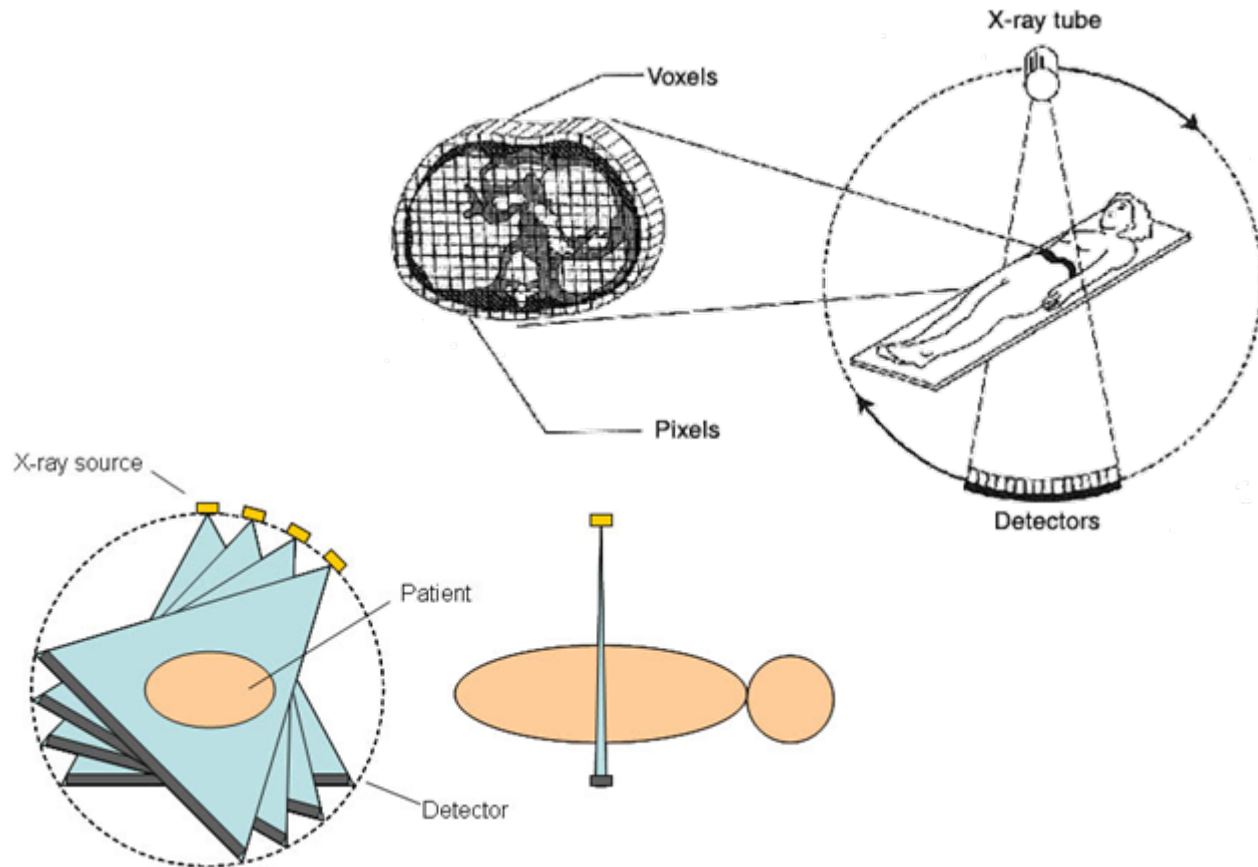
DATA: X-ray imaging in multiple plane -- by how much do x-rays get attenuated?

MAPPING_FCT: wave propagation and attenuation (optics, geometrical ray approximation)



Inverse problems are common (also outside geophysics)

Medical Imaging: Computed Tomography



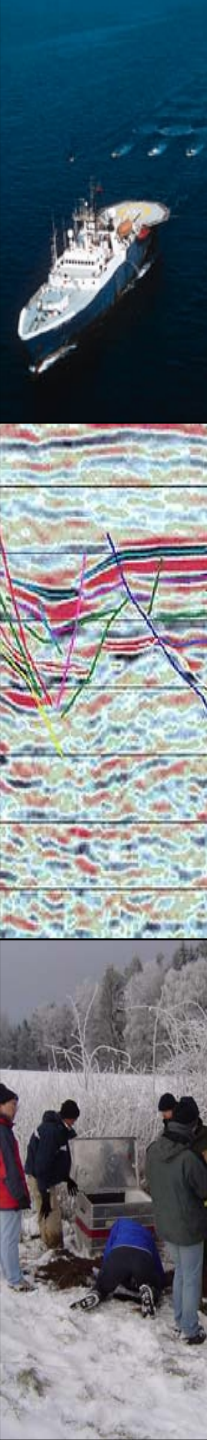
Inverse problems are common

Planetary Science: Composition of a Jupiter moon

EARTH_PROPERTY: density as function of x, y, z
DATA: gravity measurements: deflection of a satellite upon its fly-by
MAPPING_FCT: Newton's law of gravity

Common theme: **measure interior properties from the outside...**

...but not all inverse problems are like that...



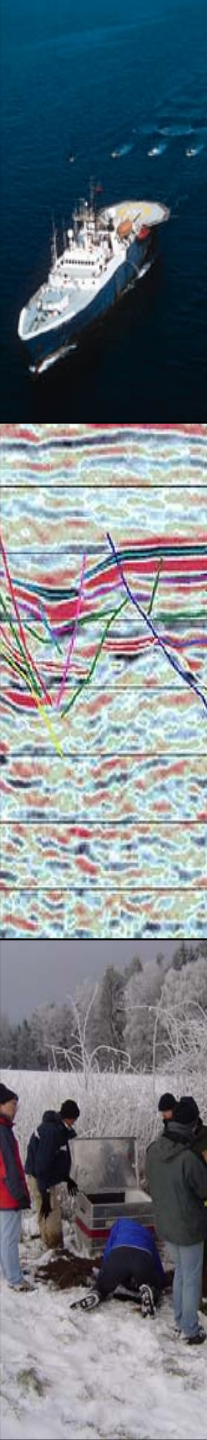
Inverse problems are common

Borehole seismics:

EARTH_PROPERTY: shallow earth properties
(velocity, density, attenuation,...) as a function of
depth

DATA: hydrophone recordings inside the borehole

MAPPING_FCT: wave equation: reflections



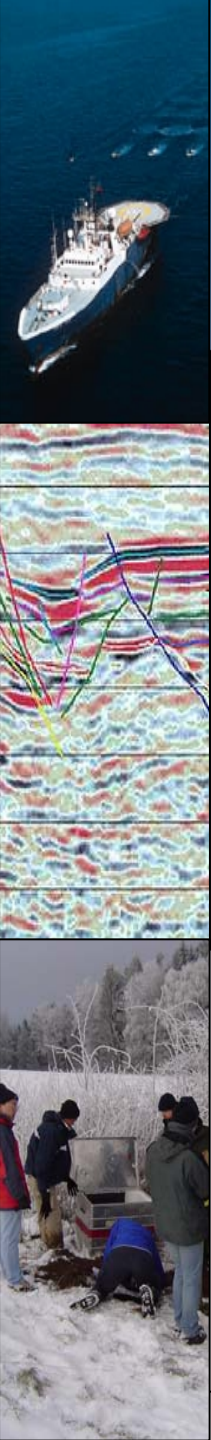
Inverse problems are common

Environmental remediation/hydrology:

EARTH_PROPERTY: source location(s) and quantity of contaminants

DATA: contaminant sensors in several deep holes around a chemical factory

MAPPING_FCT: diffusion equation/transport in porous media



Summary: What is an inverse problem?

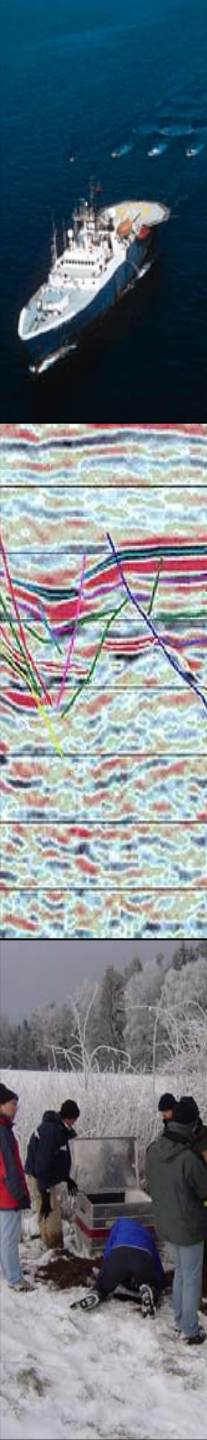
We are unable to directly measure an interesting **EARTH_PROPERTY**.

Instead we measure some other **DATA**, because we know how to derive/compute a physical relationship **MAPPING_FCT** so that:

$$\text{DATA} = \text{MAPPING_FCT}(\text{EARTH_PROPERTY})$$

We try to find the „inverse“ **MAPPING_FCT⁻¹**, so that

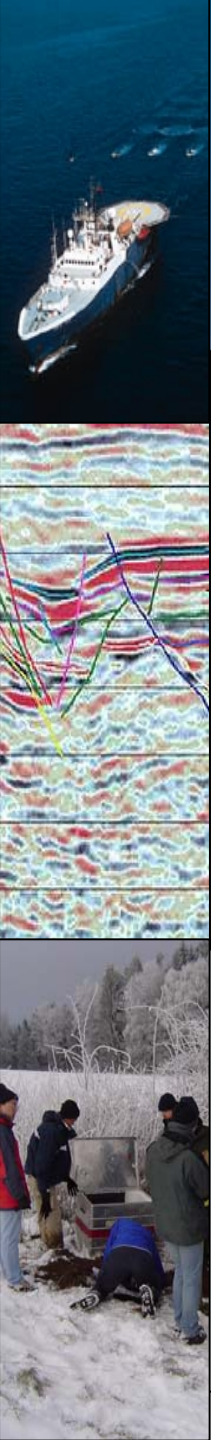
$$\text{EARTH_PROPERTY} = \text{MAPPING_FCT}^{-1}(\text{DATA})$$



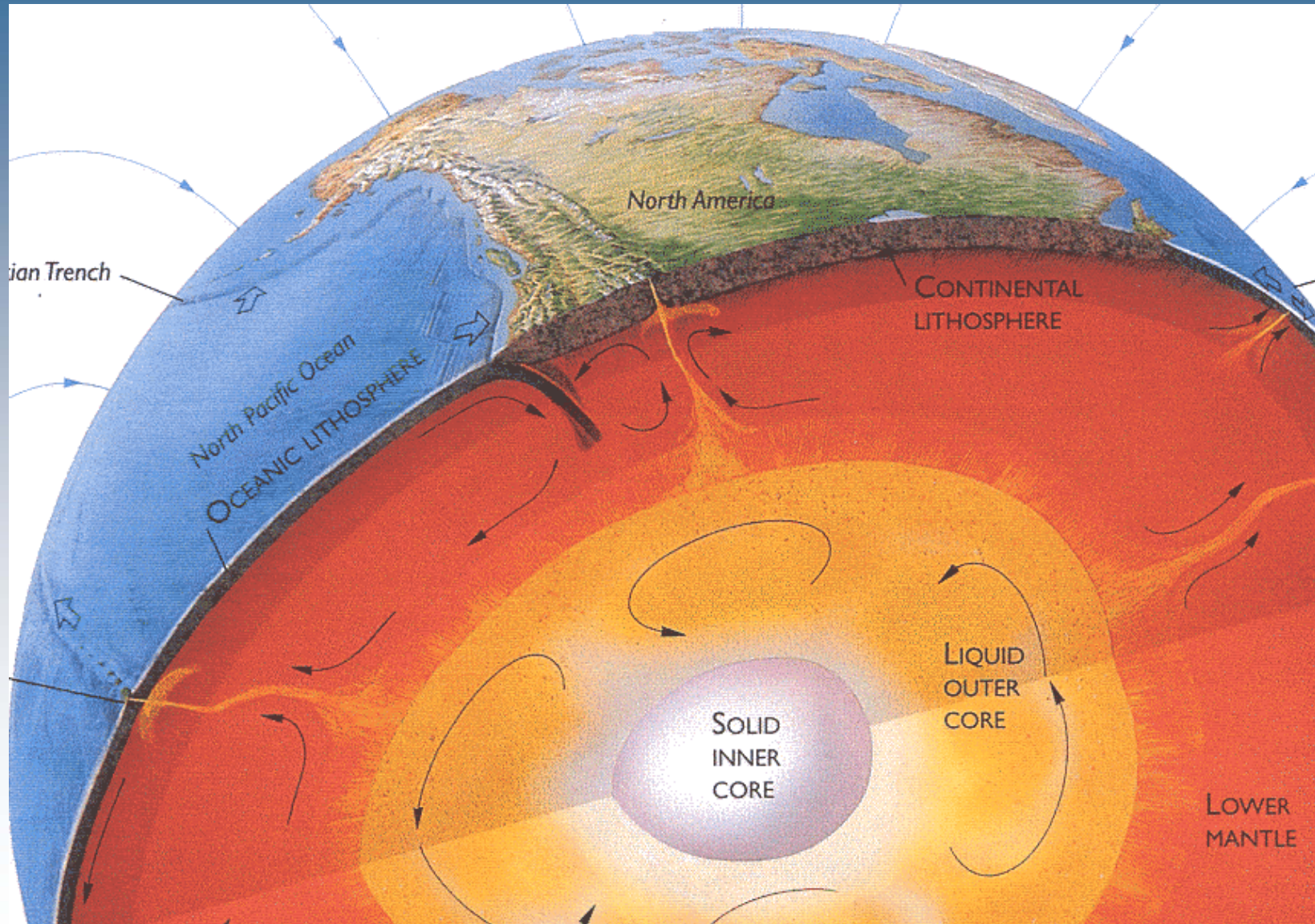
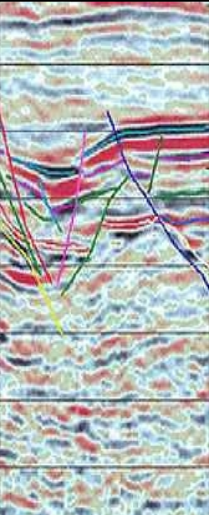
A realistic experiment: Seismic tomography of the Earth's mantle

Geophysicists' mission: Discover new things about the Earth.

If it is a good problem then many other people (geologists, geodynamicists, economists, etc.) will be interested in the results.

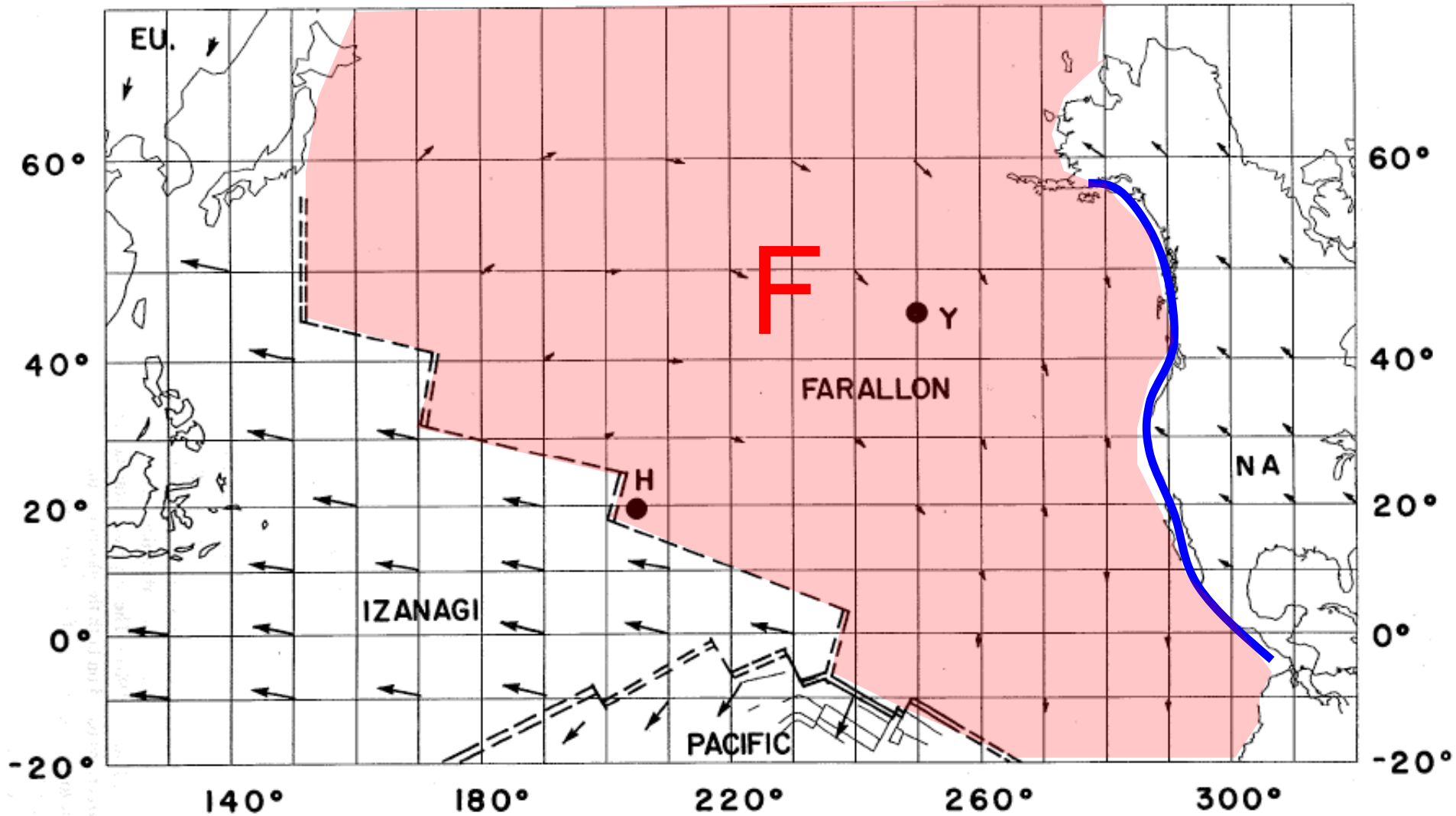


Imaging the subducted Farallon plate under North America



The Farallon plate 140 Myr ago...

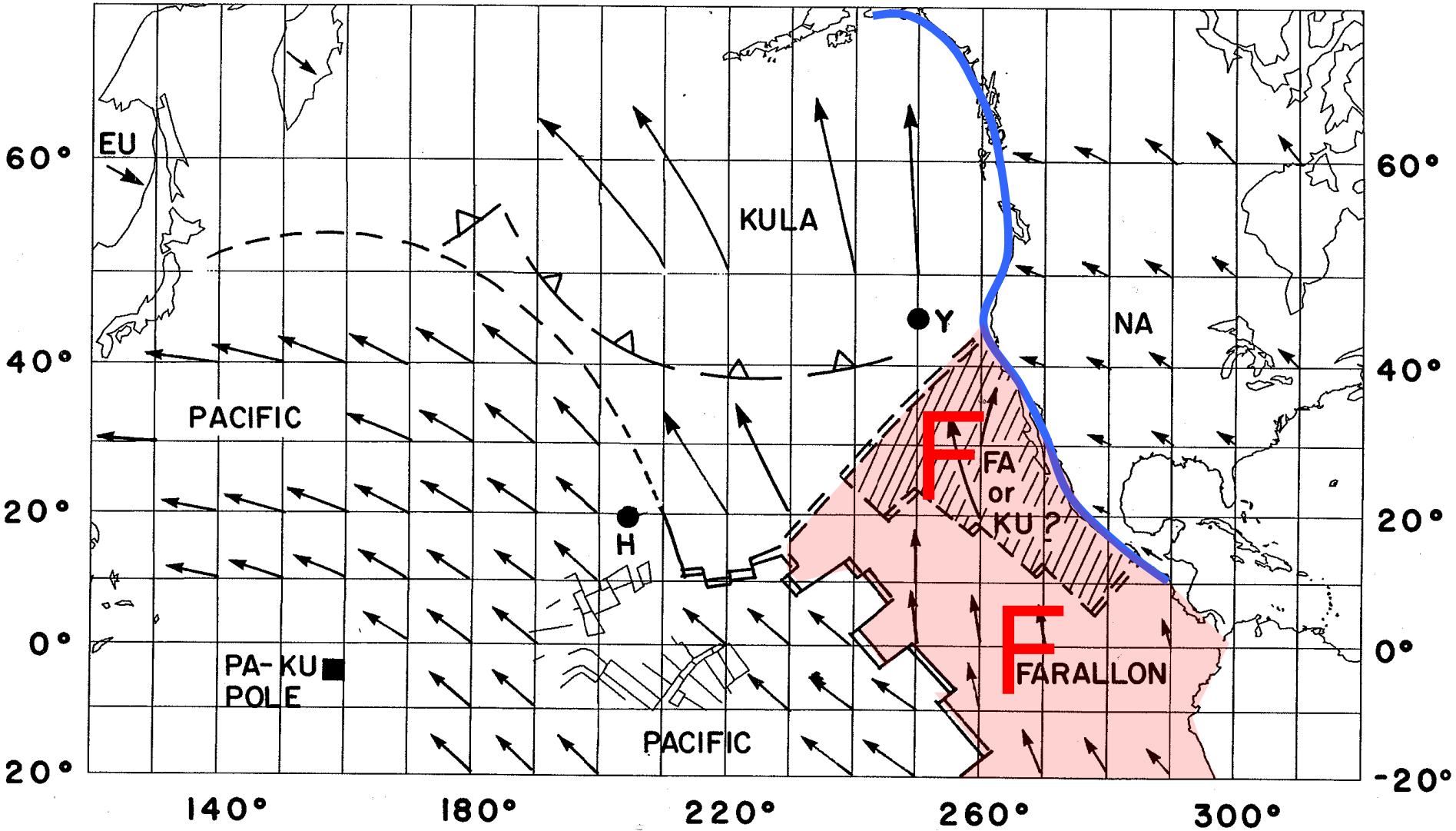
140 Ma (135-145)



Engelbreton et al. 1985

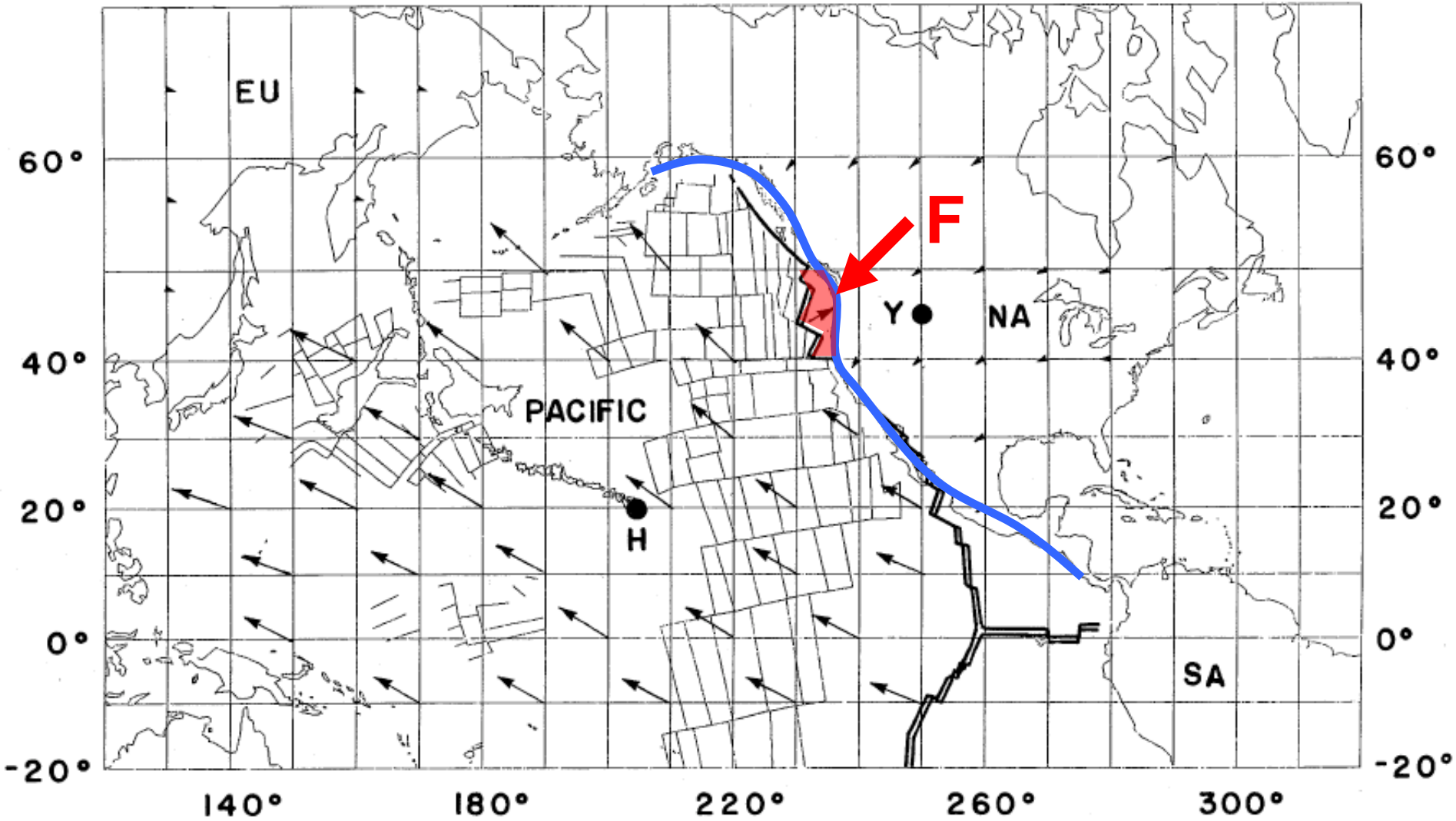
...and 80 Myr ago...

80 Ma (74-85)



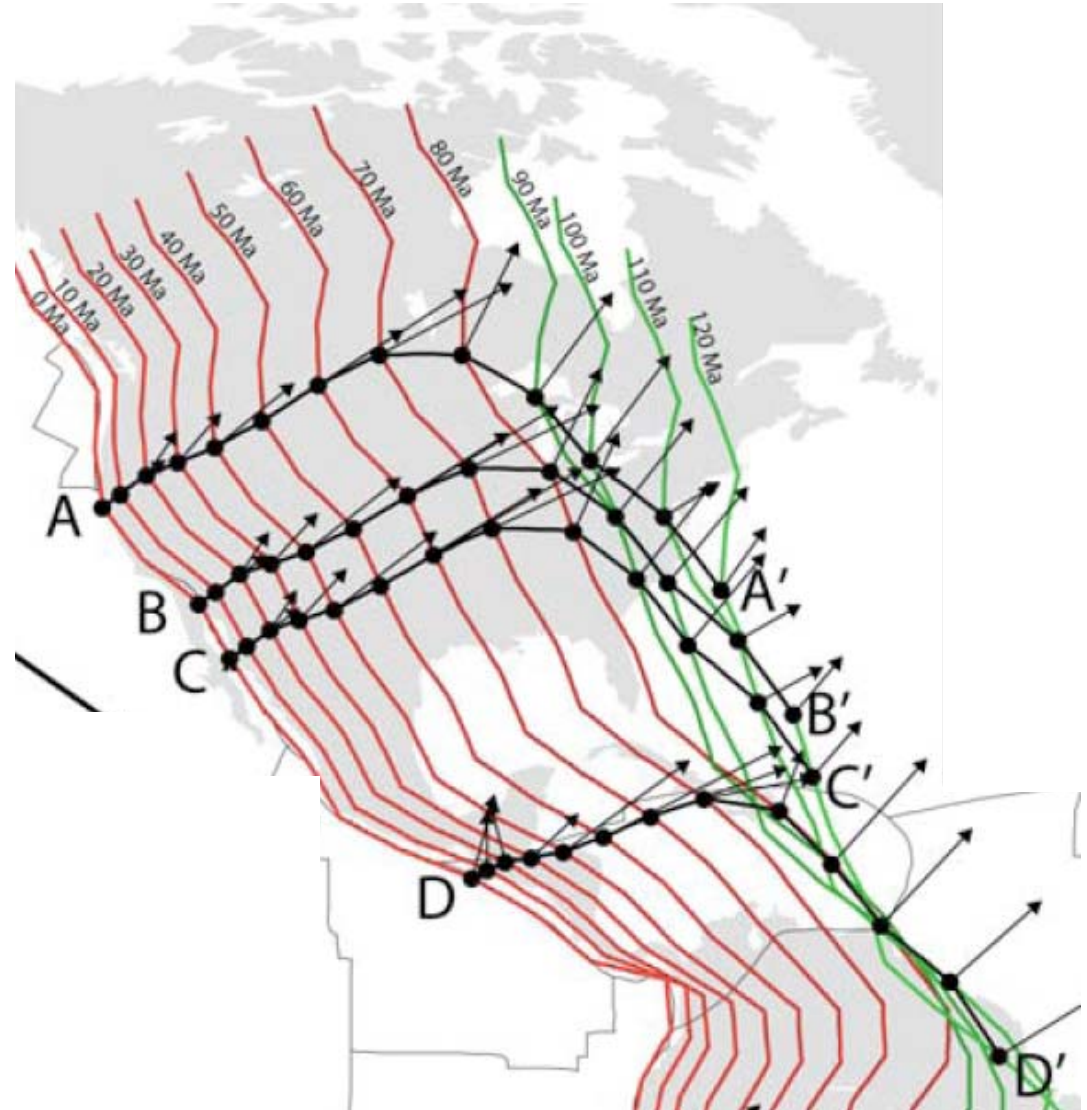
...and today.

PRESENT (0-5)



150 million years of textbook-like subduction?

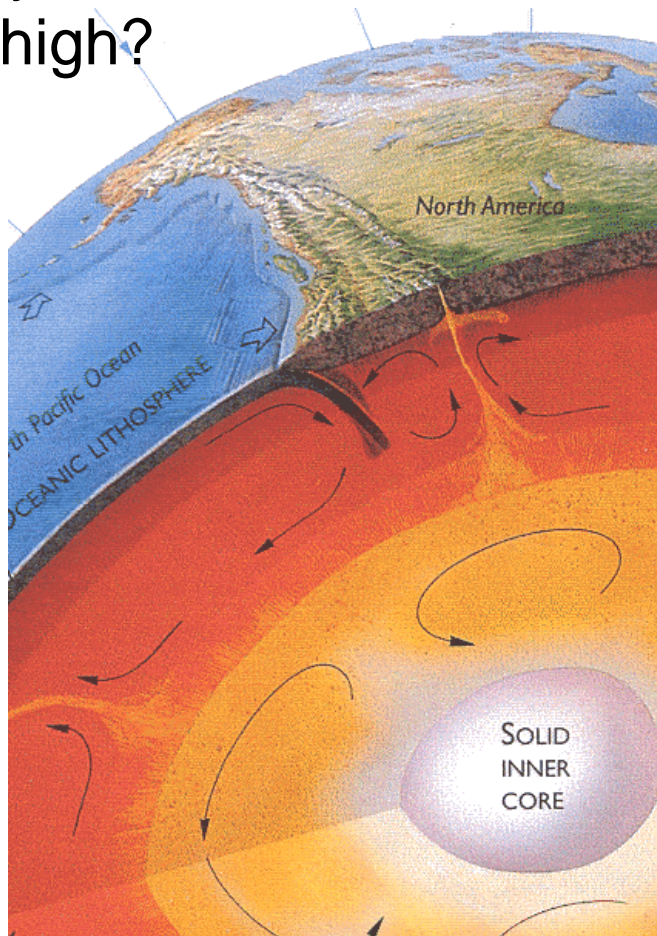
A single large plate has been subducting beneath the North American west coast for 150+ million years. No significant interference from other plates.



A simple story? Yes, but.

Extensive mountain building and volcanism far inland (since ~70 Myr). Not a “conventional” volcanic arc.

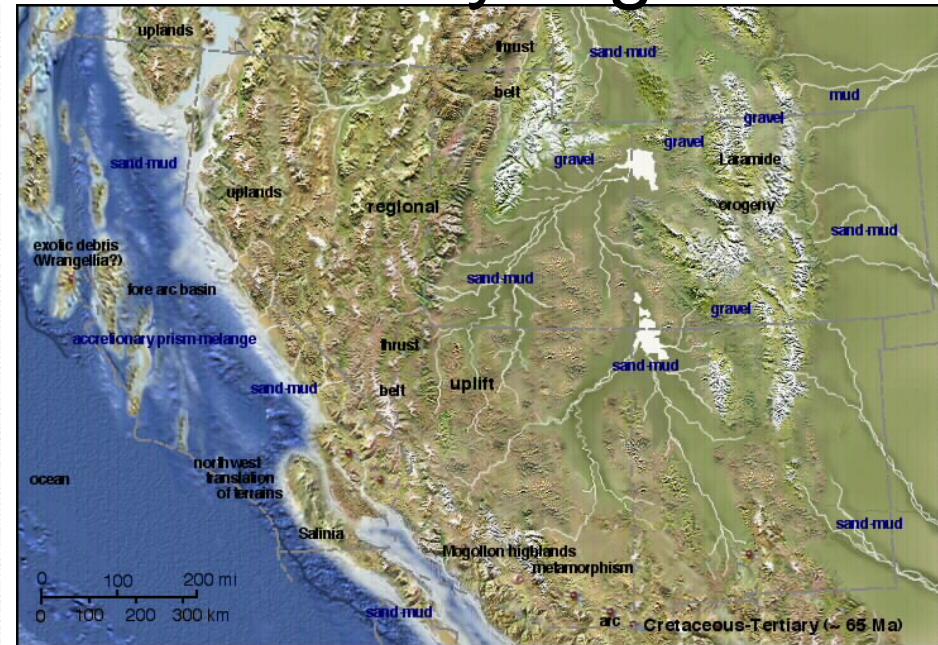
Why is the North American Cordillera so wide and stands so high?



The “Laramide orogeny”: Rapid uplift, far inland at ~70 Myr ago

75 Myr ago

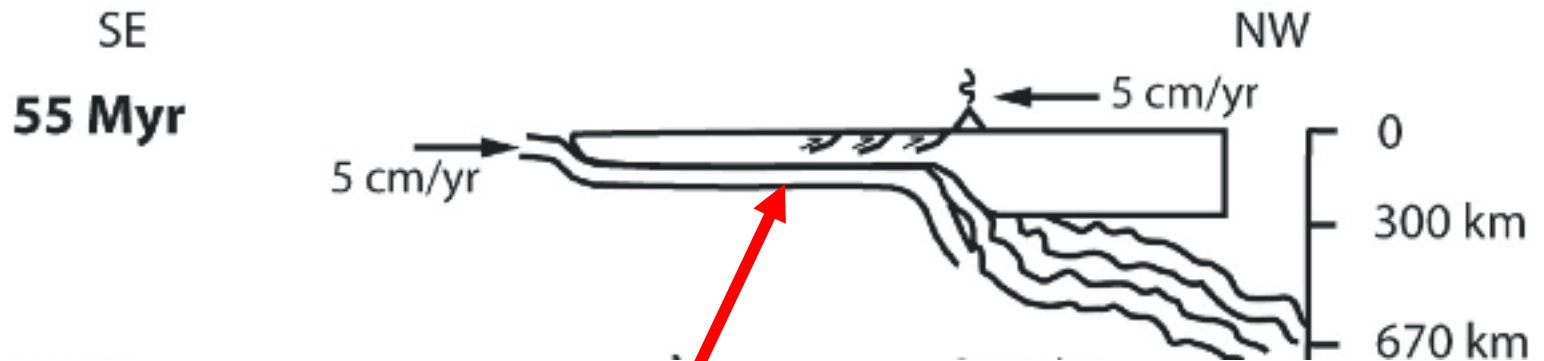
65 Myr ago



A shallow inland sea covers the Rocky Mountain area

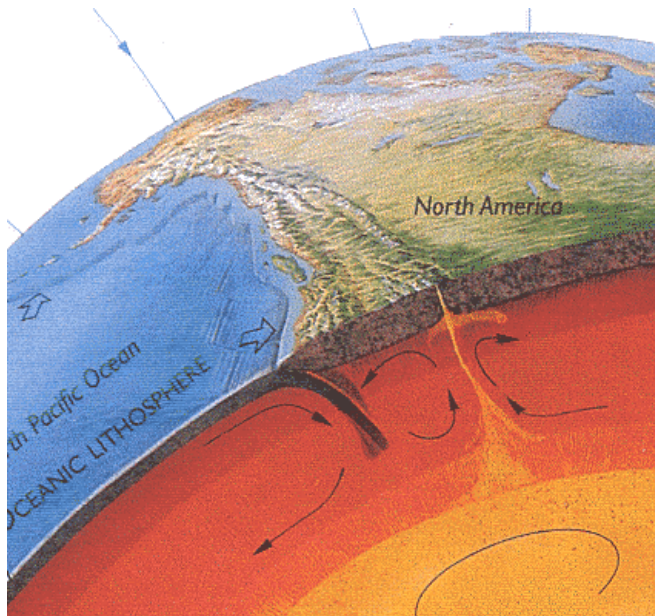
Laramide orogeny (70-50 Myr):
basement uplift by thrust faulting,
volcanic arc along trench has shut off.

Geologists' explanation: Laramide thrust faulting was caused by anomalously flat subduction



Extremely flat slab scrapes along bottom of continental crust

...but Western North America has stood high ever since.

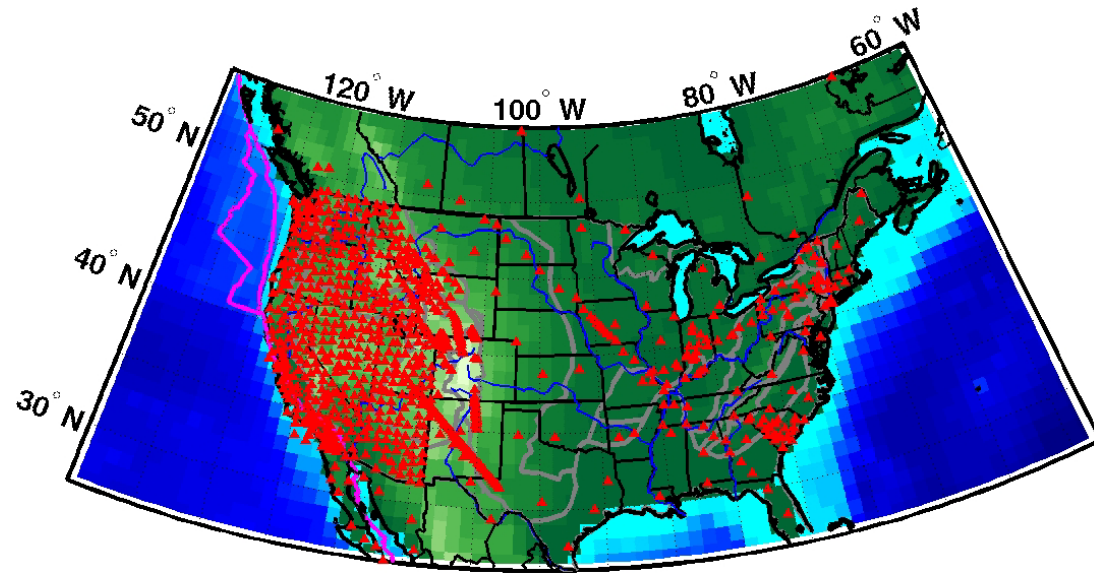
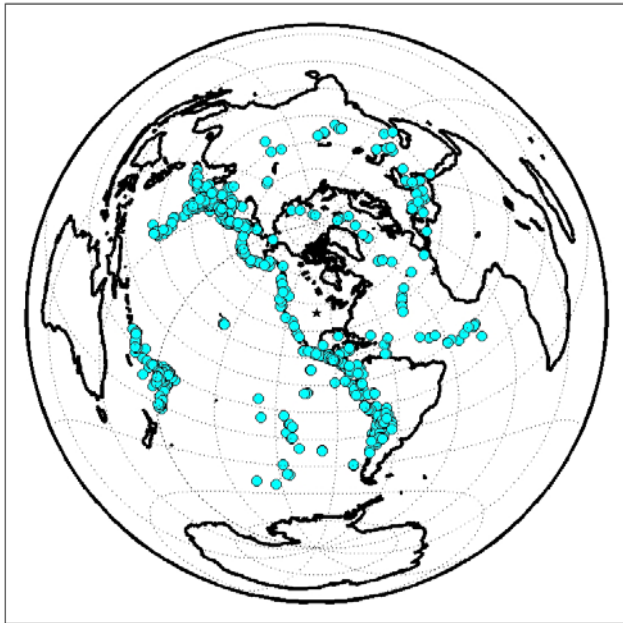


NASA satellite photo of Western U.S.; mountains are from Laramide times

Our tomographic experiment

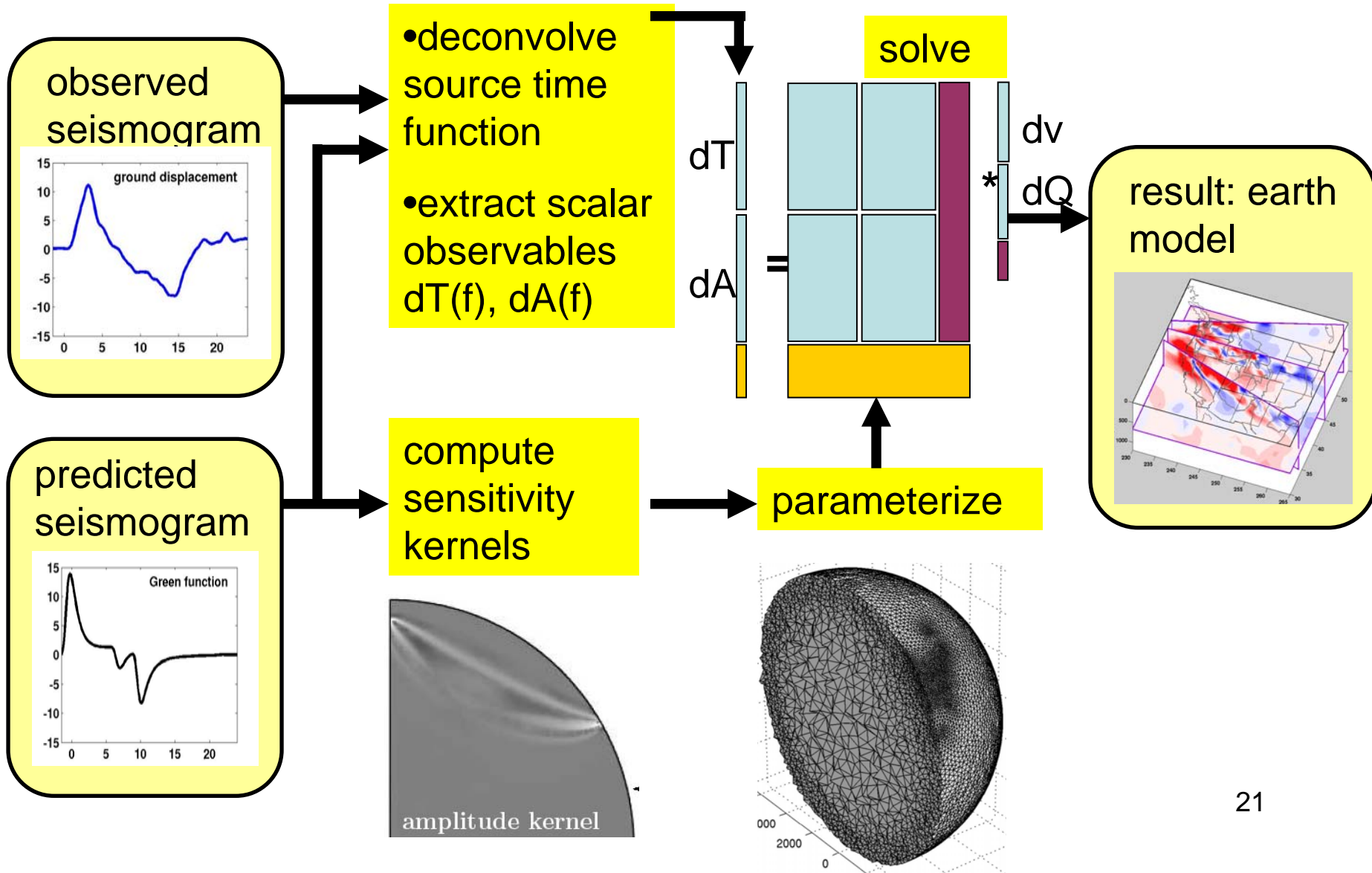
637 earthquake sources

1125 broadband receivers (seismometers)



- We use teleseismic P-wave seismograms from large earthquakes (magnitude ≥ 5.8 , 1990-2007)
- Many new USArray stations in Western U.S since 2005

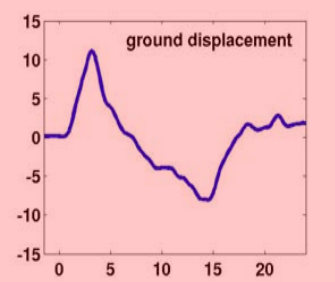
Tomography step by step



Tomography step by step

DATA

observed
seismogram



- deconvolve source time function
- extract scalar observables $dT(f)$, $dA(f)$

dT

dA

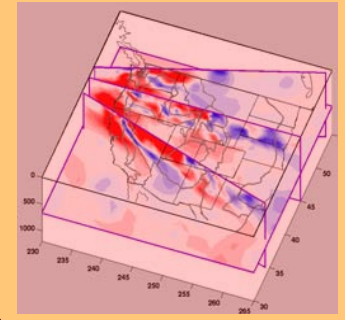
solve

EARTH PROPERTY

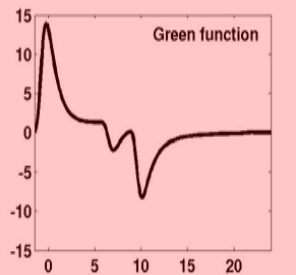
dv

dQ

result: earth
model

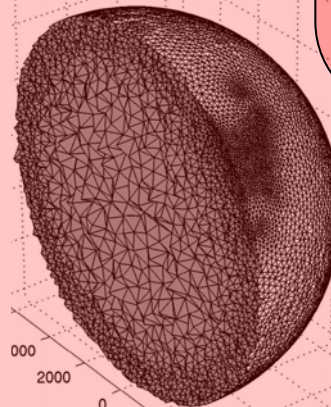


predicted
seismogram



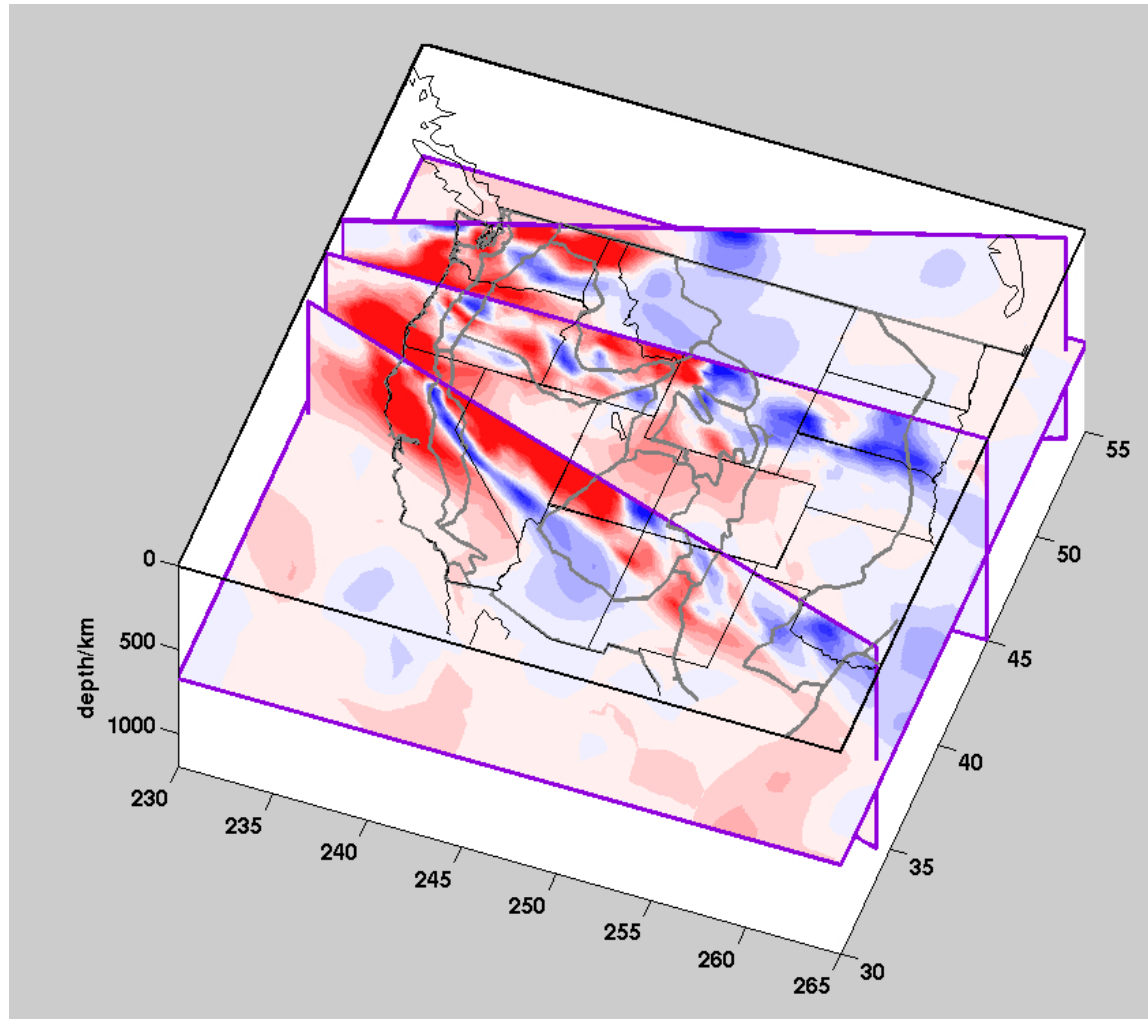
compute
sensitivity
kernels

parameterize

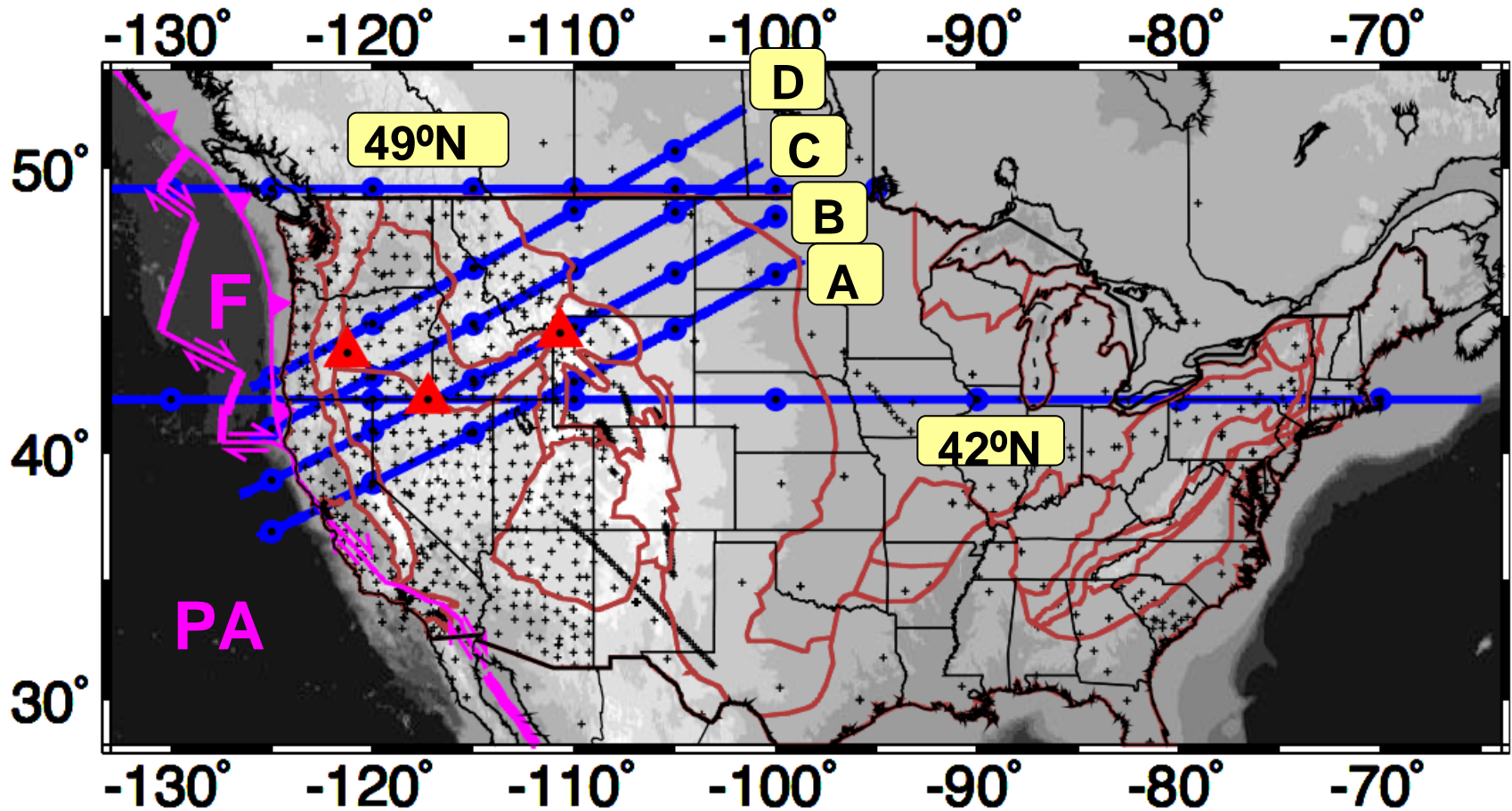


MAPPING_FCT

Result: 3-D model of P-wave velocities under North America



Locations of interesting cross-sections



The big picture: Not one, but two episodes of whole mantle subduction

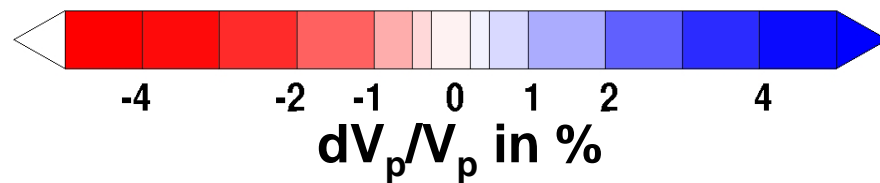
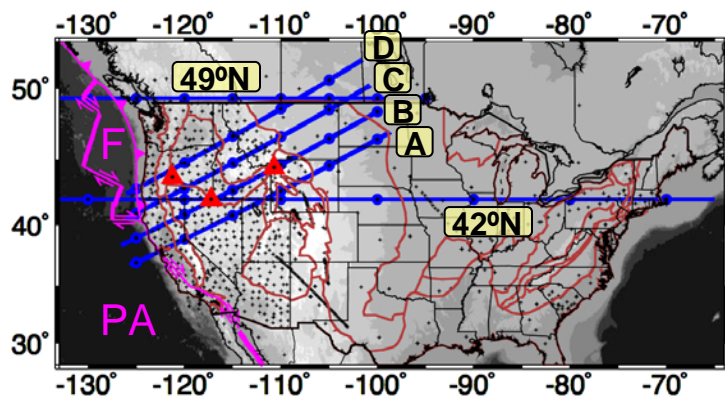
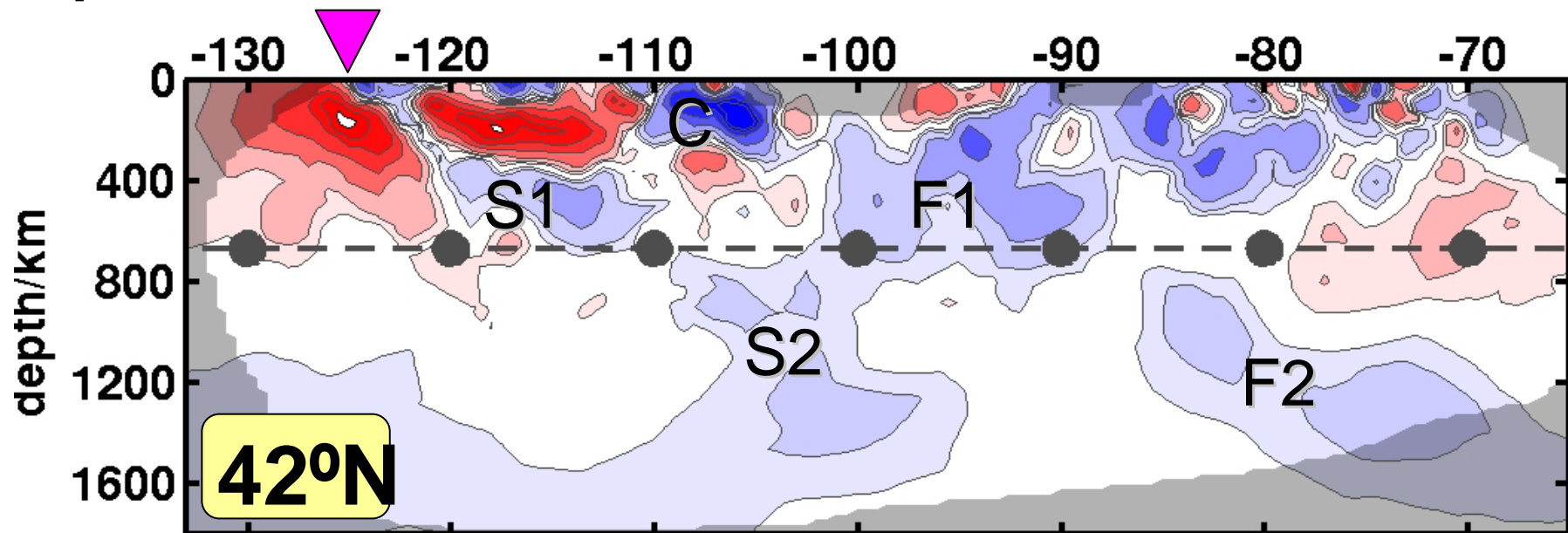
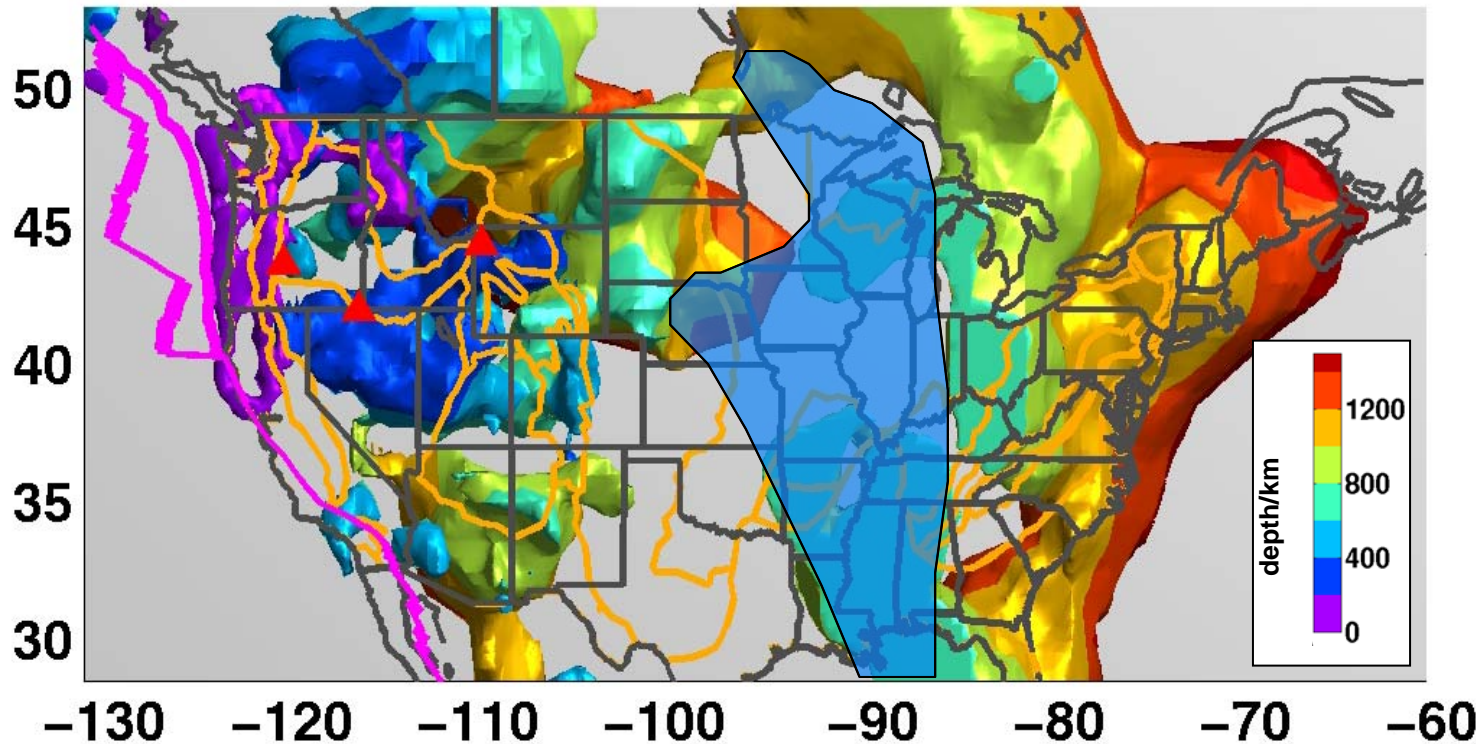
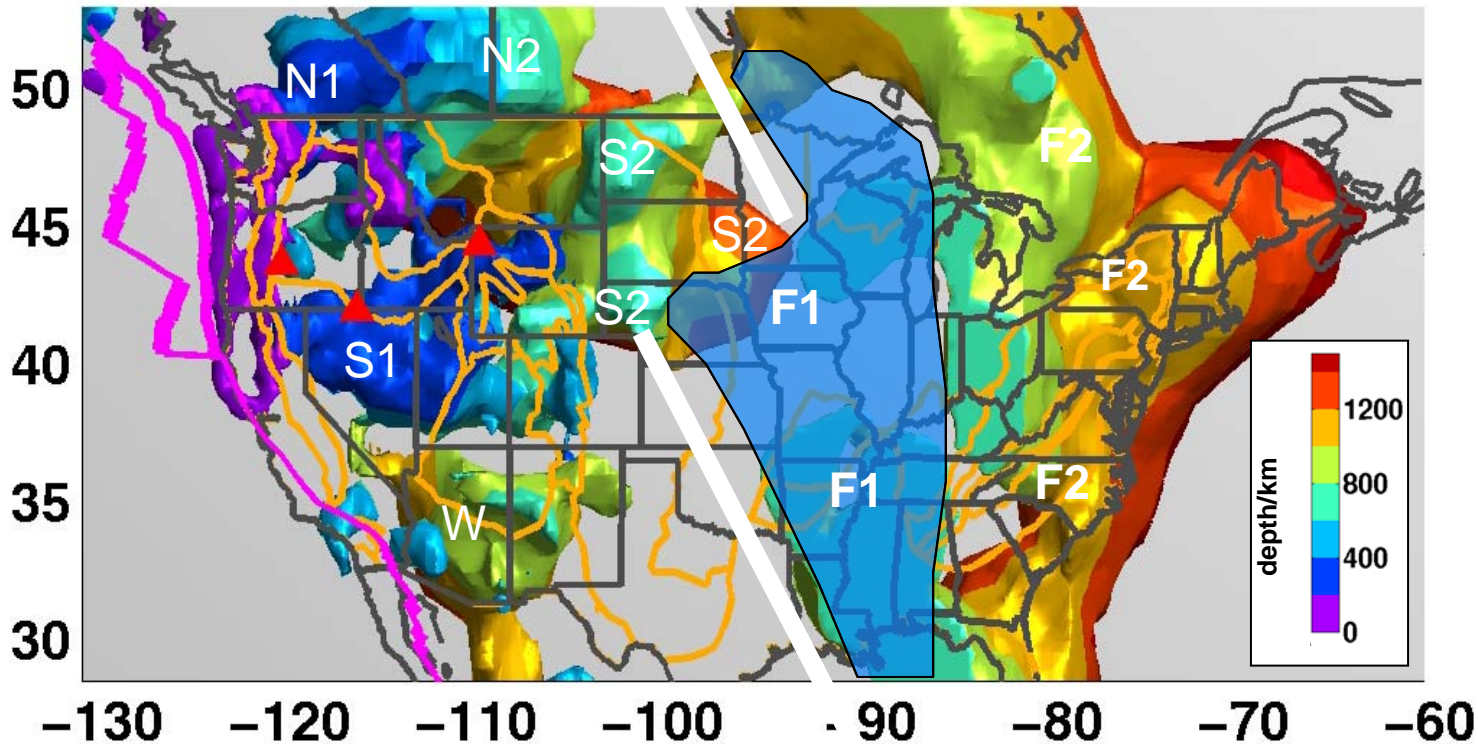


Image of the subducted Farallon slab in the mantle



- Seismically fast material is contoured (fast means cold).
- Color signifies depth. We can confidently image ~1500 km deep.
- Crust and lithosphere not rendered.

Interpretation: a frontal plate break ended the Laramide era at ~50 Myr



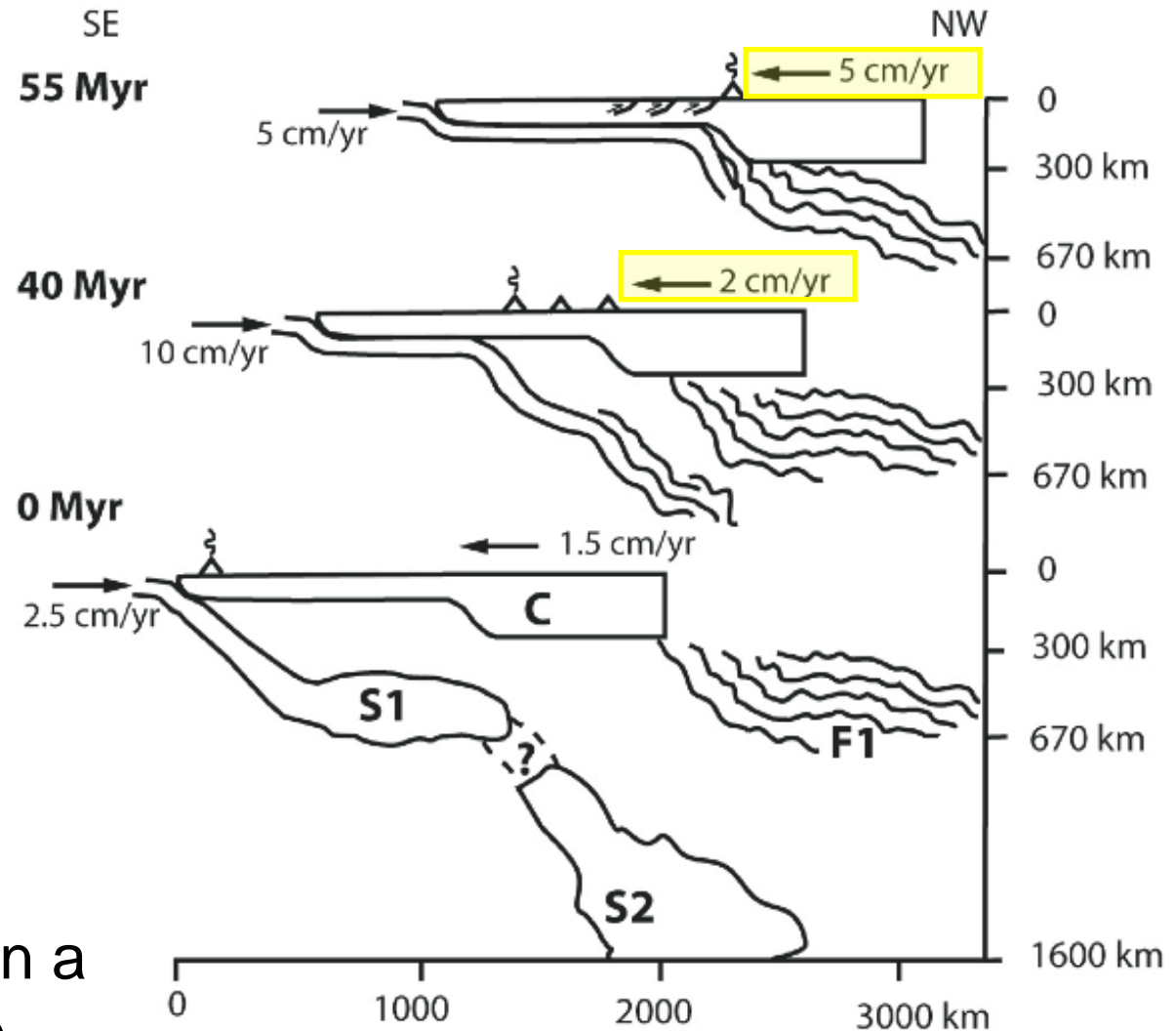
- F1 must have been part of the Laramide flat slab. It still fills the transition zone.
- Lower end of S2/N2 subducted ~55 Myr ago.

Interpretation: a frontal plate break ended the Laramide era at ~50 Myr

55 Myr ago

40 Myr ago

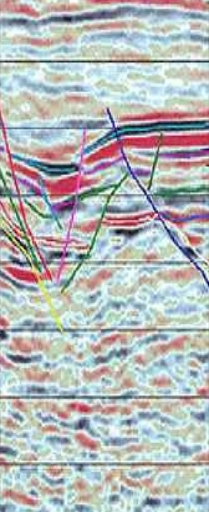
Today



All velocities in a hotspot frame.

How does this work?

Some intuitive examples...



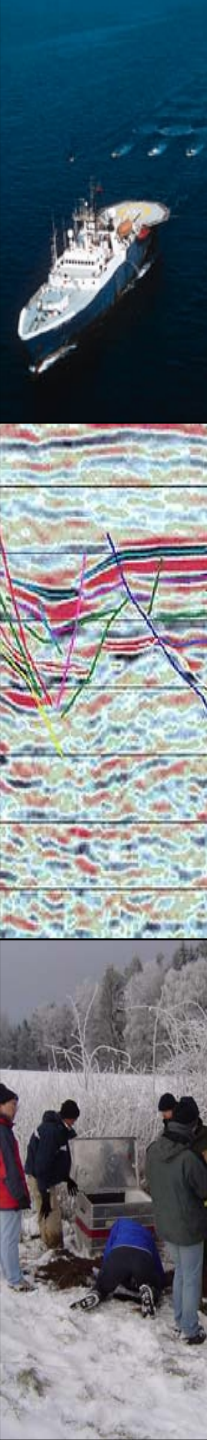
Tomography: Intuitive example 1

Surface waves (of a certain frequency) have sampled the shallow mantle of North America, along the shown raypaths.

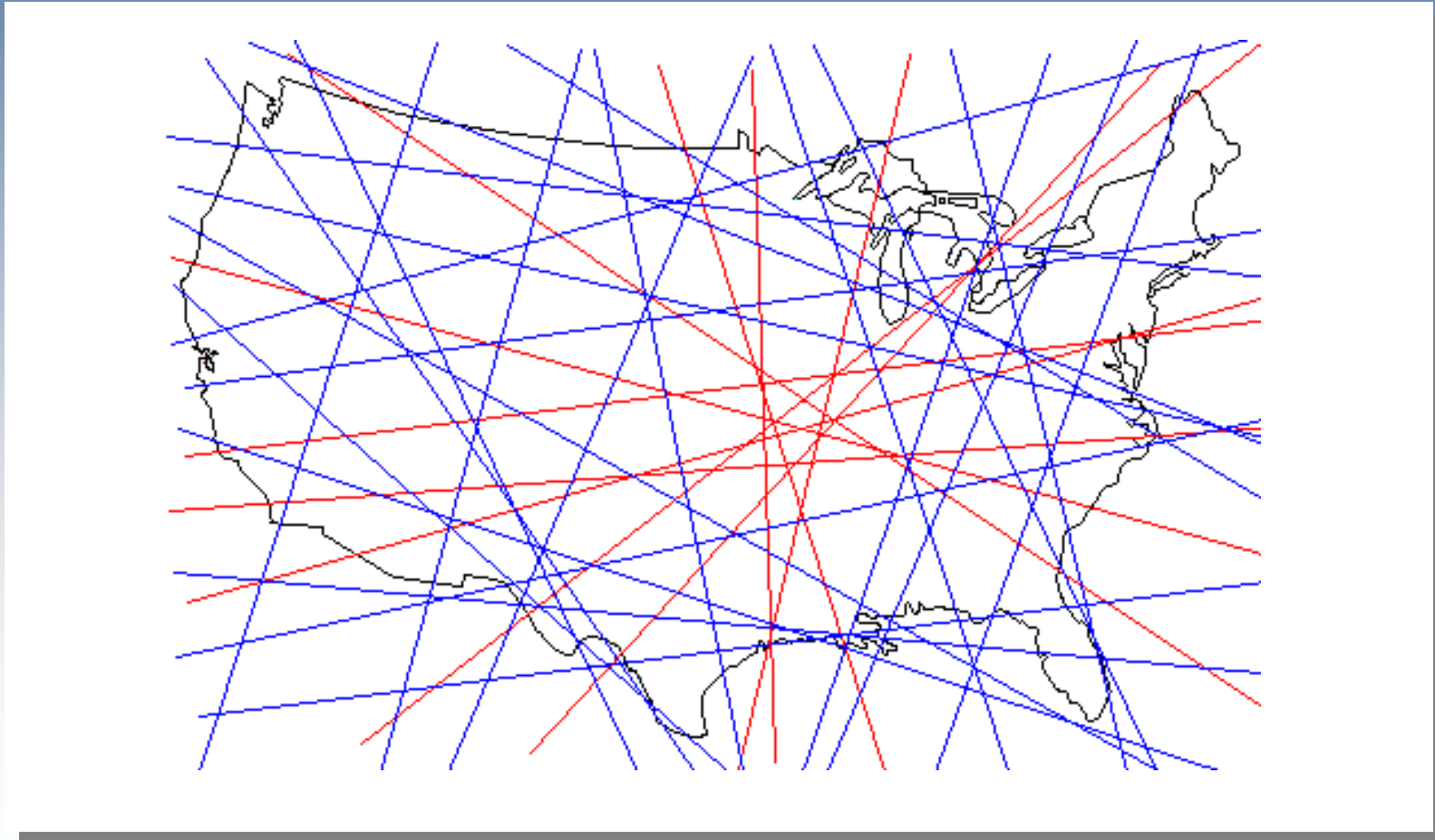
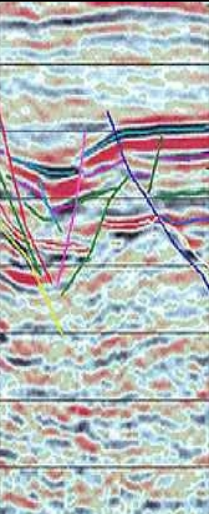
Your prior guess is that traveltimes \sim length of ray, meaning $v(x,y) = v_0 = \text{const. everywhere}$.

In reality you observe anomalies in the traveltimes DATA:
Red ray means traveltimes were longer than expected.
Blue ray means traveltimes were as expected.

Red rays must have traversed some slow material. Where exactly is it located?

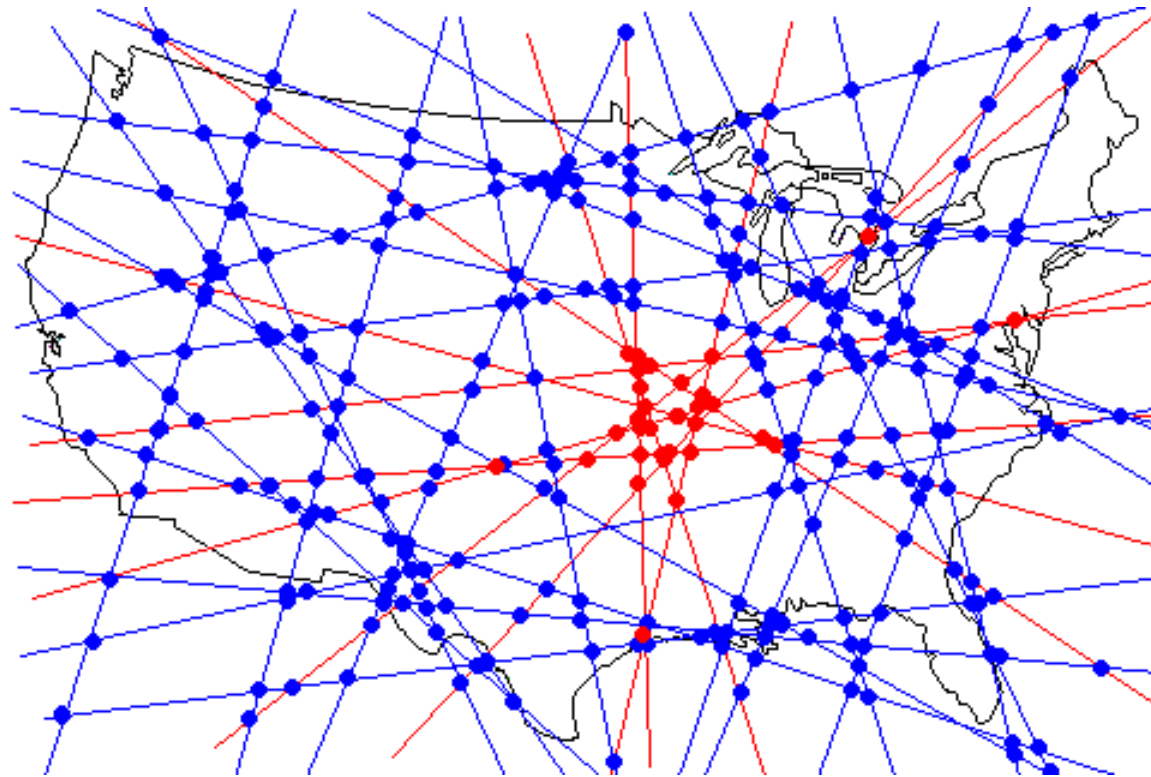


Tomography: Intuitive example 1



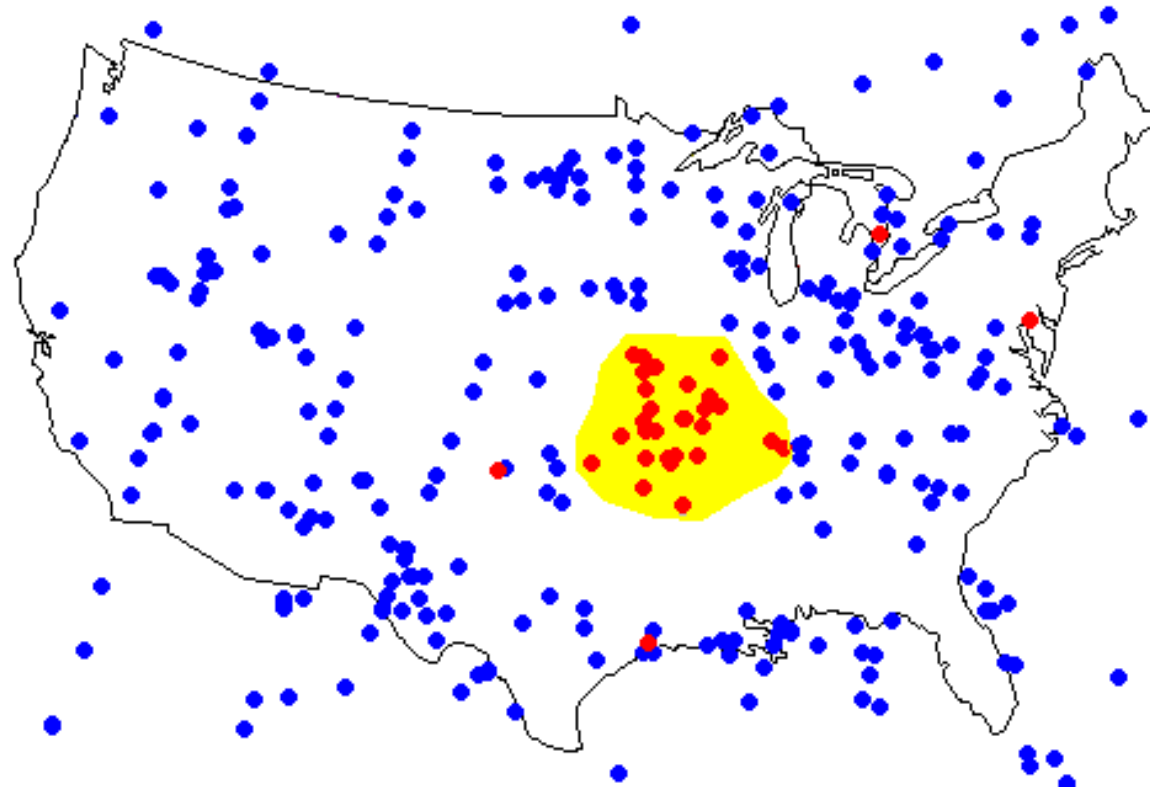
Tomography: Intuitive example 1

Idea: Consider ray crossings



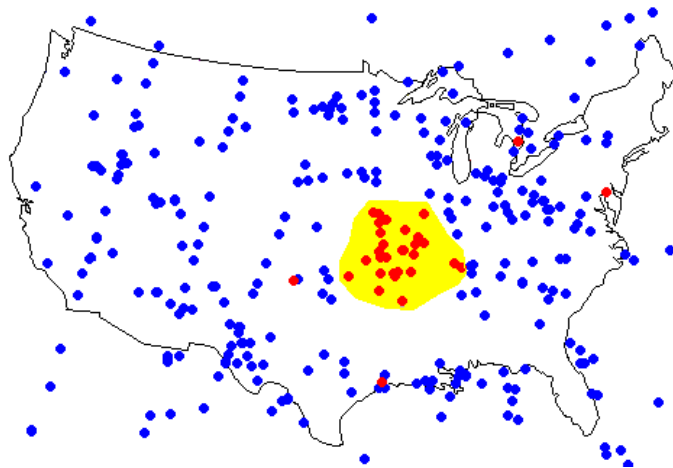
Tomography: Intuitive example 1

Idea: Consider ray crossings

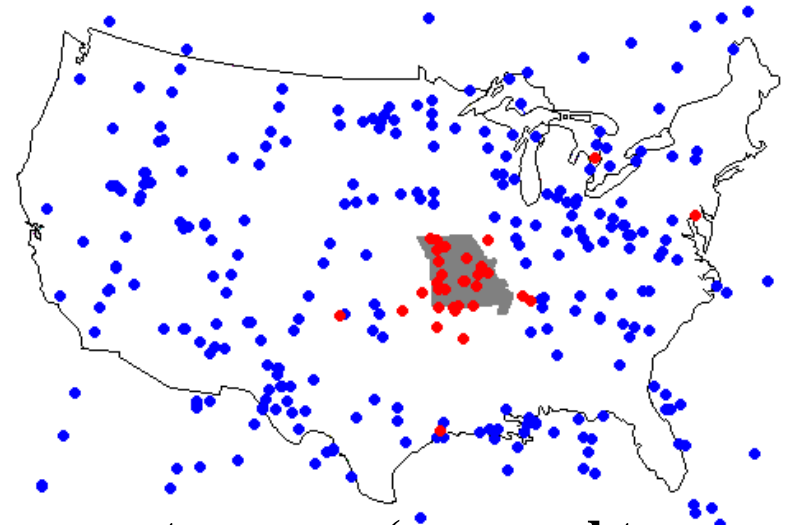


Tomography: Intuitive example 1

It worked pretty well.



recovered area

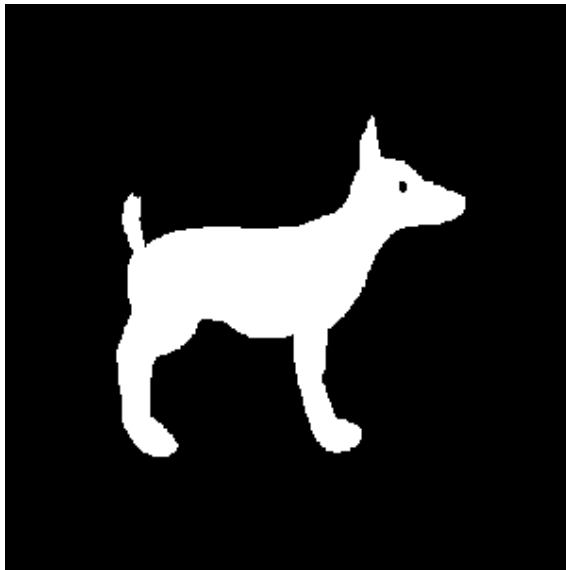


true area (was used to generate the colored rays)

Why is the reconstruction not perfect?

Tomography: Intuitive example 2

Image reconstruction



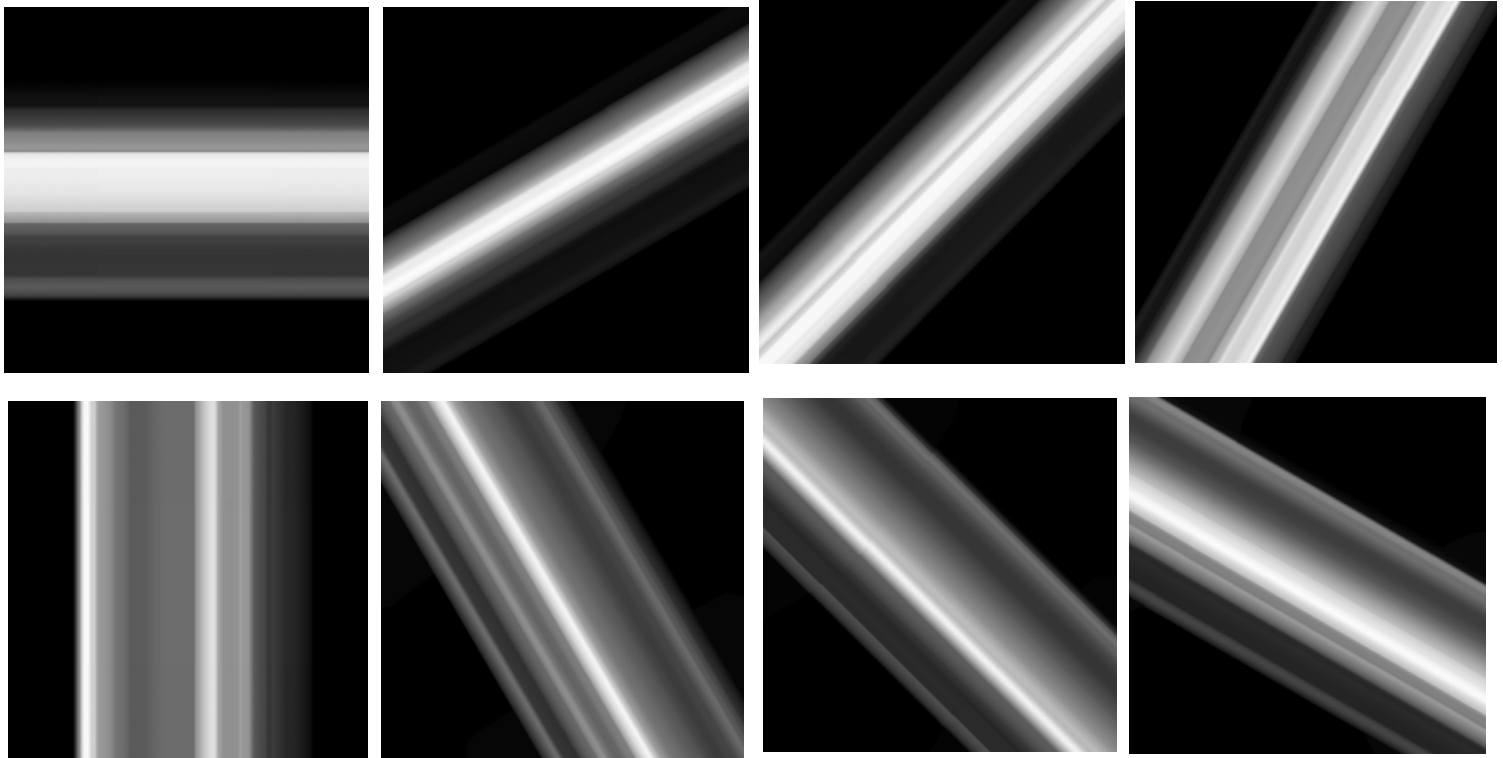
original image



We smear the image in horizontal direction (like an x-ray integrates over different body tissues along its path)

Tomography: Intuitive example 2

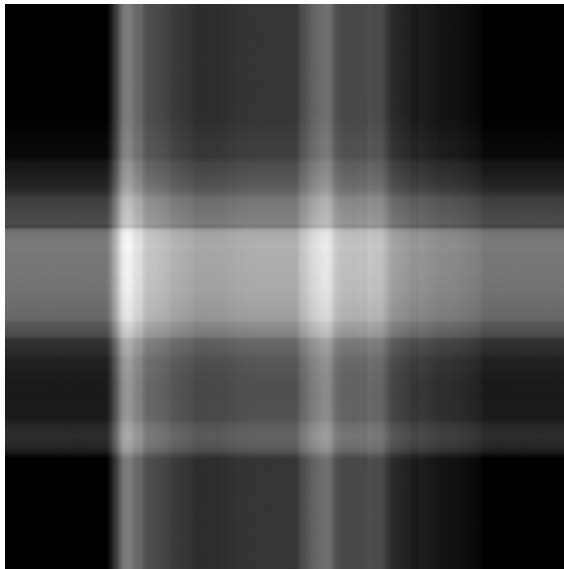
Image reconstruction: Generating „DATA“



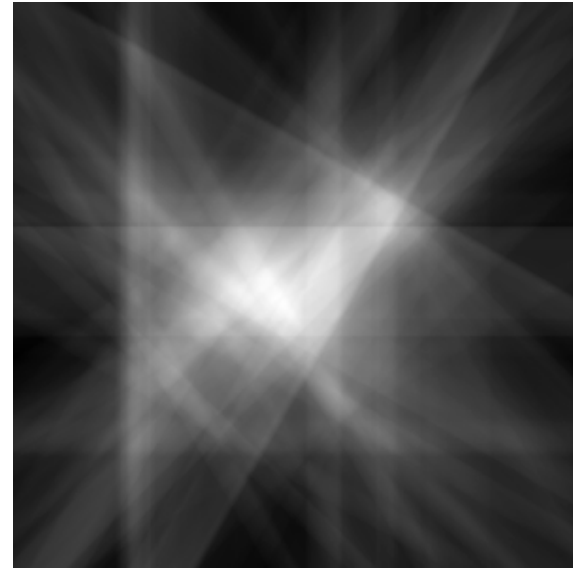
**We smear over more directions to simulate more x-ray “data”:
what rays “see” from all these different angles**

Tomography: Intuitive example 2

Now we try to reconstruct the image (principle of destructive/constructive interference):



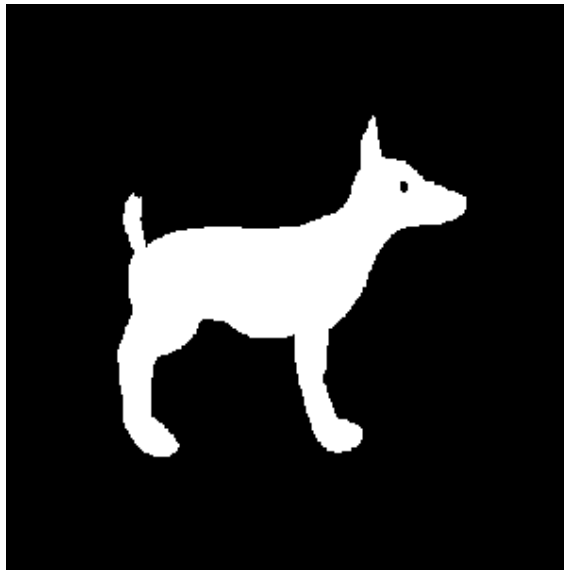
**Addition of two
directions of the “data”**



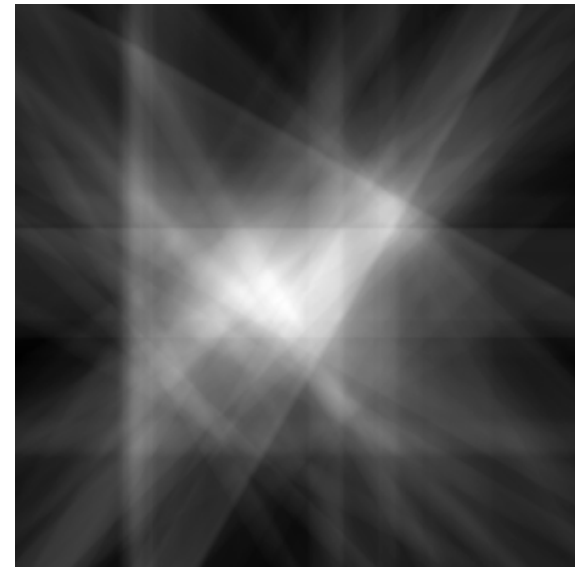
**Addition of all 8
directions of the “data”**

Tomography: Intuitive example 2

Reconstruction result



Original



Reconstruction

How could we improve on this?

Basic modeling

Acoustic tomography:

A few bricks are standing next to each other. To first order they all have the same, known P-velocity v_0 (or slowness $u_0 = 1/v_0$), except for small variations: $u_i = u_0 + \Delta u_i$, where $\Delta u_i \ll u_0$. We want to estimate the small anomalies Δu_i .

$$u_0 + \Delta u_1$$

$$u_0 + \Delta u_2$$

$$u_0 + \Delta u_3$$

Basic modeling

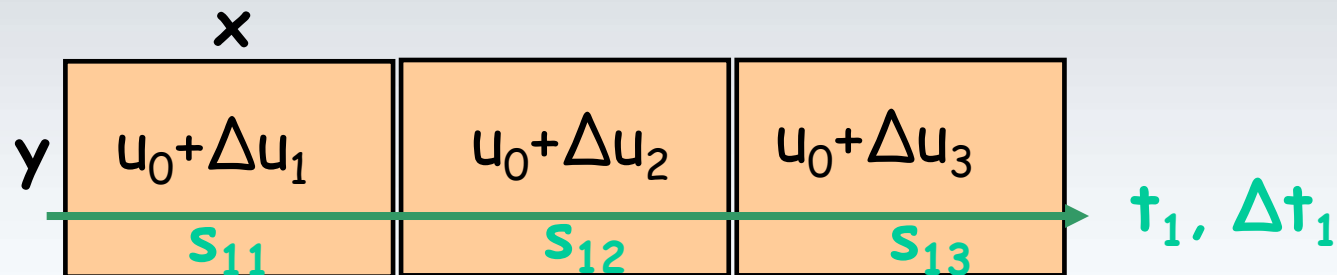
Acoustic tomography:

Blocks are x wide and y high.

Traveltime: $t_1 = u_1 s_{11} + u_2 s_{12} + u_3 s_{13}$

Traveltime anomaly: $\Delta t_1 = \Delta u_1 s_{11} + \Delta u_2 s_{12} + \Delta u_3 s_{13}$

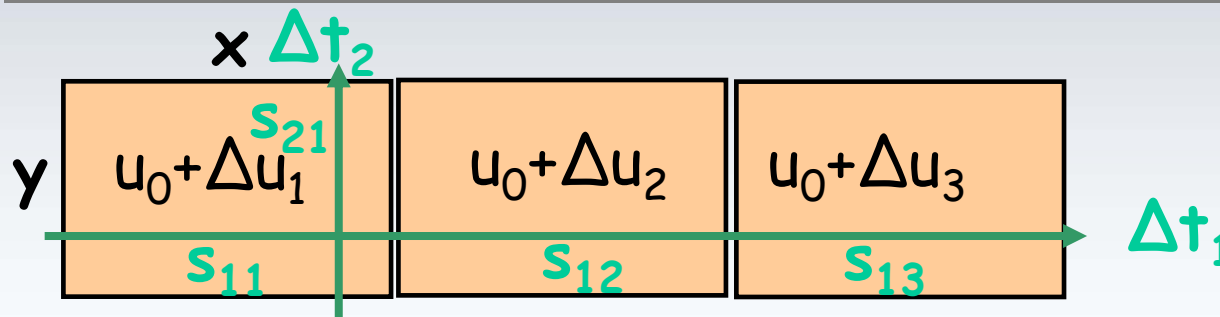
The s_{ij} can be computed from the given geometry
(general case = Snell's law!)



Basic modeling

Acoustic tomography:

Linear system: two equations, three unknowns
Matrix notation:



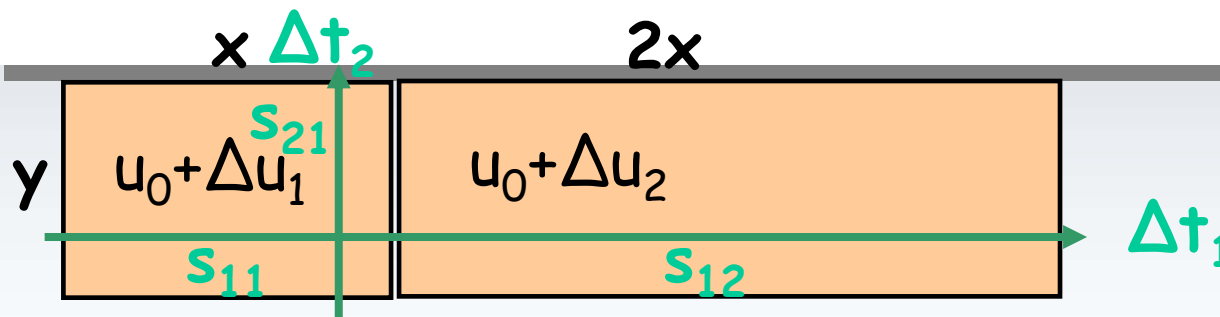
Basic modeling

Acoustic tomography:

For full rank: two equations, two unknowns.

Full rank means:

Matrix notation (and its inverse):

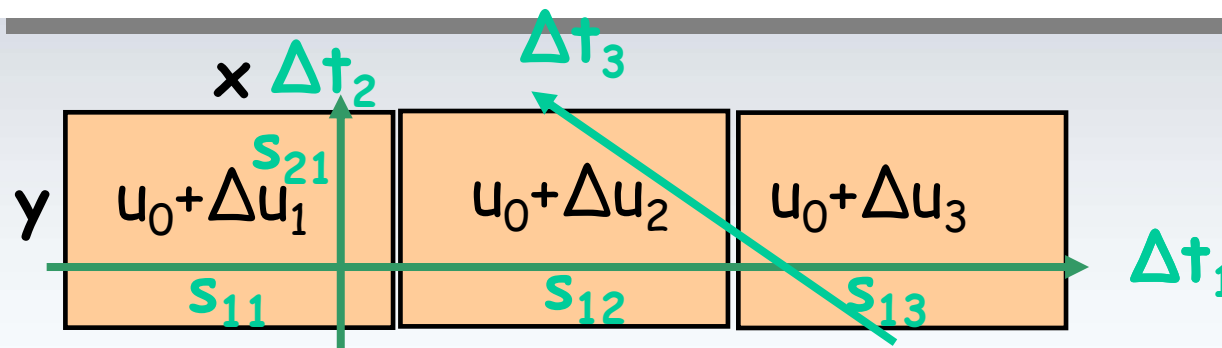


Basic modeling

Acoustic tomography:

Does this system have full rank?

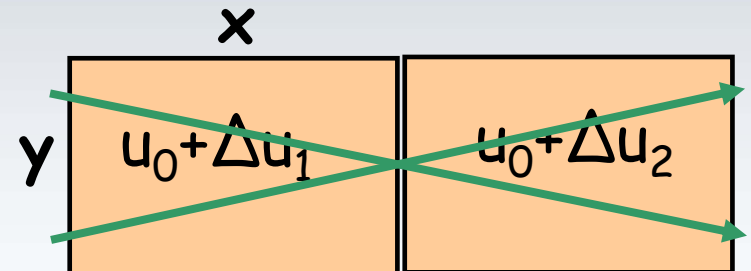
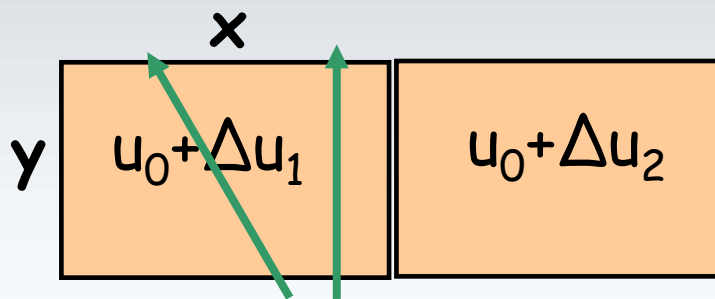
How many measurements M need to be made for the matrix to have an inverse?



Basic modeling

Acoustic tomography: How about these geometries?

Ideally, each measurement should contribute as much new information as possible („independent“ measurements --> experiment design)



Real-world experiments

Parameterizations of the unknowns (grid)



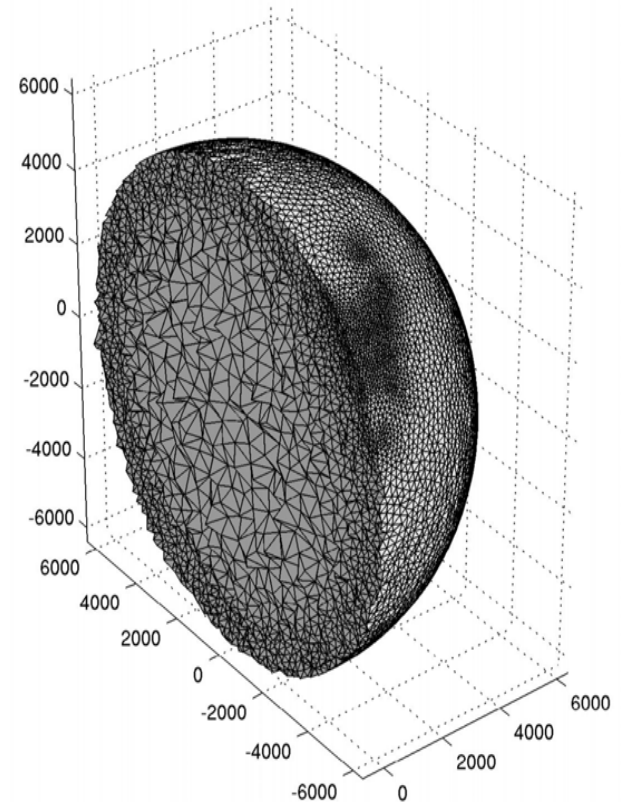
Δu_1

Δu_2

Coarse parameterization in blocks; few unknowns

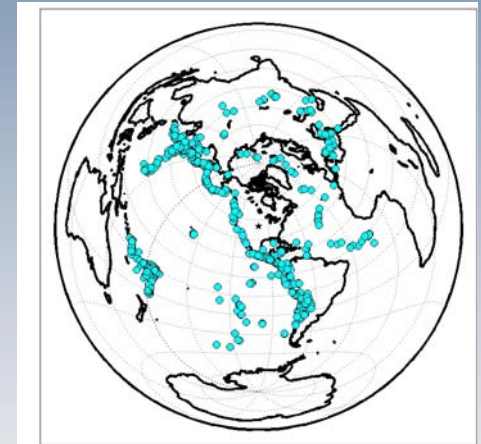
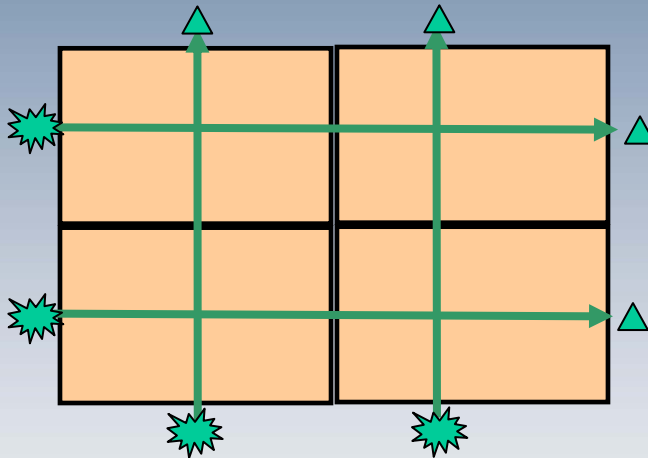
vs.

Complex parameterization (irregular tetrahedra, 10^5 unknowns)



Real-world experiments

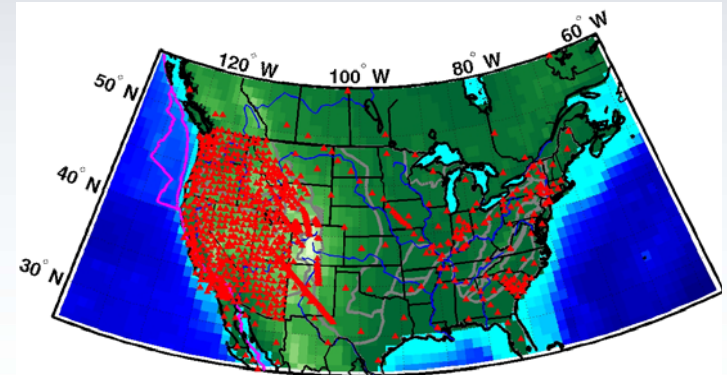
Source and receiver geometry



Optimally designed source-receiver geometry

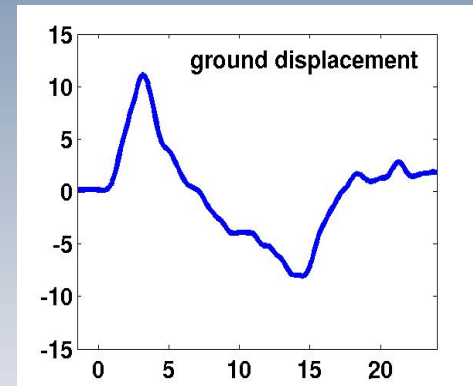
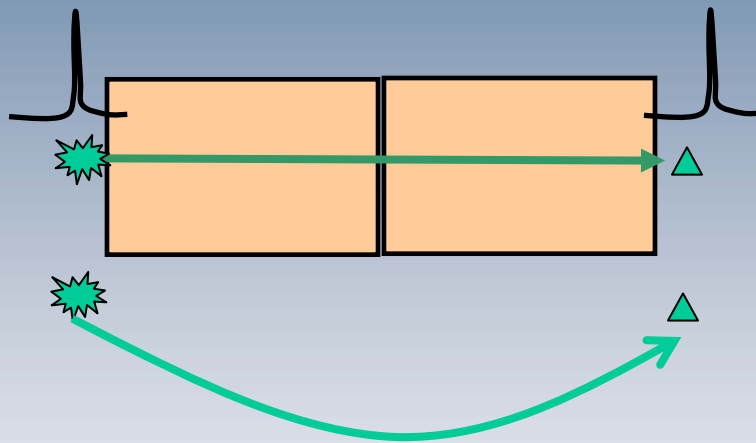
vs.

"Take what you can get" -->



Real-world experiments

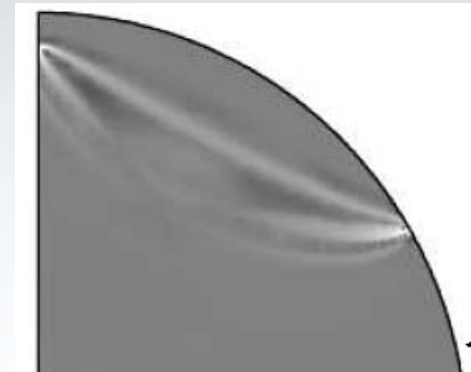
Signals and wave propagation modeling



Sharp pulses modelled as
optical rays

vs.

Realistic wavelets with broad
Fresnel zones -->



Real-world experiments

System to solve

Small system, well conditioned,
exactly determined

vs.

Huge system, ill conditioned, both
underdetermined and
overdetermined

ART OF TOMOGRAPHY:
Finding smart ways to
solve this anyway.

